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Key Points:

- We investigated the limitations of biogeochemical models in simulating inter-annual variability (IAV) of gross primary production (GPP)
- Capturing IAV of model parameters and diurnal GPP peaks can be key to understanding IAV of GPP
- Variability in model performance is jointly influenced by model types, parameterization strategies, and site characteristics

Supporting Information:

Supporting Information may be found in the online version of this article.

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Addressing Challenges in Simulating Inter–Annual Variability of Gross Primary Production

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Abstract A long-standing challenge in studying the global carbon cycle has been understanding the factors controlling inter–annual variation (IAV) of carbon fluxes, and improving their representations in existing biogeochemical models. Here, we compared an optimality-based model and a semi-empirical light use efficiency model to understand how current models can be improved to simulate IAV of gross primary production (GPP). Both models simulated hourly GPP and were parameterized for (a) each site–year, (b) each site with an additional constraint on IAV ($Cost^{IAV}$), (c) each site, (d) each plant–functional type, and (e) globally. This was followed by forward runs using calibrated parameters, and model evaluations using Nash–Sutcliffe efficiency (NSE) as a model-fitness measure at different temporal scales across 198 eddy-covariance sites representing diverse climate–vegetation types. Both models simulated hourly GPP better (median normalized NSE: 0.83 and 0.85) than annual GPP (median normalized NSE: 0.54 and 0.63) for most sites. Specifically, the optimality-based model substantially improved from NSE of -1.39 to 0.92 when drought stress was explicitly included. Most of the variability in model performances was due to model types and parameterization strategies. The semi-empirical model produced statistically better hourly simulations than the optimality-based model, and site–year parameterization yielded better annual model performance. Annual



Georg Wohlfahrt, Ankur R. Desai, Eugénie Paul-Limoges, Marta Galvagno, Albin Hammerle, Georg Jocher, Borja Ruiz Reverter, David Holl, Jiquan Chen, Luca Vitale, M. Altaf Arain Formal analysis: Ranit De Methodology: Ranit De, Shanning Bao, Sujan Koirala, Nuno Carvalhais Resources: Markus Reichstein, Nuno Carvalhais Software: Ranit De Supervision: Sujan Koirala, Alexander Brenning, Markus Reichstein, Nuno Carvalhais Validation: Ranit De Visualization: Ranit De Writing - original draft: Ranit De Writing - review & editing: Ranit De. Shanning Bao, Sujan Koirala, Alexander Brenning, Markus Reichstein, Torbern Tagesson, Michael Liddell, Andreas Ibrom, Sebastian Wolf, Ladislav Šigut, Lukas Hörtnagl, William Woodgate, Mika Korkiakoski, Lutz Merbold, T. Andrew Black, Marilyn Roland, Anne Klosterhalfen, Peter D. Blanken, Sara Knox, Simone Sabbatini, Bert Gielen, Leonardo Montagnani, Rasmus Fensholt, Georg Wohlfahrt, Ankur R. Desai, Eugénie Paul-Limoges, Marta Galvagno, Albin Hammerle, Georg Jocher, Borja Ruiz Reverter, David Holl, Jiquan Chen, Luca Vitale, M. Altaf Arain, Nuno Carvalhais

model performance did not improve even when parameterized using $Cost^{IAV}$. Furthermore, both models underestimated the peaks of diurnal GPP, suggesting that improving predictions of peaks could produce better annual model performance. Our findings reveal current modeling deficiencies in representing IAV of carbon fluxes and guide improvements in further model development.

Plain Language Summary Terrestrial vegetation assimilates and releases carbon dioxide through photosynthesis and respiration, respectively, and their net magnitude determines if vegetation can be a sink or source of carbon dioxide. We are interested in understanding what controls the inter–annual variability (IAV) of gross primary production (GPP) which represents photosynthesis, and how their representations can be improved in models simulating GPP. Here, we considered an optimality-based model that can be applied equally well globally, and a data-driven semi-empirical model. We found both models better simulated diurnal and seasonal cycles than the IAV of GPP. Such differences probably stem from model parameters, as critical ecosystem functions they represent may not be well-constrained or model structures may lack critical representations via inaccurate simulation of peak diurnal GPP and drought stress. The IAV of GPP was comparatively better simulated if model parameters were fine-tuned with data from specific years. Another challenge is that IAV of GPP can also be observed due to disturbances, such as forest fire, and human management besides natural causes, which were also not represented in models. Our results suggest that learning the variability of model parameters over the years can be key to better simulation of the IAV of GPP.

1. Introduction

The global carbon cycle is an important biogeochemical cycle, which affects the climate on Earth (Schimel, 2001). Terrestrial vegetation, which covers a large part of the land area, assimilates atmospheric carbon dioxide (CO₂) through photosynthesis. Simultaneously, CO₂ of similar magnitude is released into the atmosphere through terrestrial ecosystem respiration (TER). The net balance of these two fluxes determines if terrestrial ecosystem acts as a sink or source of carbon (Ruehr et al., 2023). Terrestrial gross primary production (GPP) can be defined as "apparent" photosynthesis, that is, the rate at which the vegetation assimilates carbon through photosynthesis minus the loss of carbon only through photorespiration (Plummer, 2006; Wohlfahrt & Gu, 2015). GPP can be estimated directly using gas exchange measurements at the leaf and canopy scales (Jez et al., 2021), and indirectly through measurements of net ecosystem exchange (NEE) using the eddy covariance (EC) method at the ecosystem or landscape scale (D. D. Baldocchi, 2003). Though the GPP estimated using the EC method represents "apparent" photosynthesis, its magnitude can be closer to "true" photosynthesis which is the actual amount of carbon assimilated due to overestimation of daytime mitochondrial respiration in flux–partitioning algorithm (Reichstein et al., 2005; Wohlfahrt & Gu, 2015). Furthermore, a large variety of biogeochemical models have been developed to simulate and upscale carbon fluxes from local to regional or global scales to better describe the global carbon cycle (Burton et al., 2023; Dannenberg et al., 2023; Nelson et al., 2024; Xiao et al., 2014).

Biogeochemical models that simulate GPP can be of different types and complexities. On the one hand, processbased models, such as the models used in the Trends in Net Land-Atmosphere Carbon Exchange (TRENDY) project, mechanistically describe the physiological processes involved in photosynthesis or plant respiration (Sitch et al., 2015). Their ability to capture a certain process largely depends on the underlying model structure and calibration of model parameters (Anav et al., 2015). Similar, but simpler than fully mechanistic approaches are the models constructed on the concept of light use efficiency (LUE), which treat a canopy usually as one big leaf, but where the GPP is calculated as the product of the absorbed photosynthetically active radiation (aPAR) and LUE (Monteith, 1972). These models are semi-empirical as they combine both the simplicity of empirical models and the theoretical mechanisms that underpin process-based models (J. Chen, 2021; Running et al., 2000; Yuan et al., 2007). On the other hand, data–driven empirical models (Jung et al., 2011, 2020) are largely based on learning regression functions to establish a general relation between input data, such as meteorology and ecosystem properties, and the desired output, such as GPP. These models' performance largely relies on good quality training data and generally lacks comprehensive representations of long-term forcing functions.

Considering the methodological diversity and differences in GPP estimates, various model benchmarking and model-data integration experiments have been designed to compare approaches and to unveil drivers of ecosystem functioning for various climate-vegetation types, across spatial and temporal scales. A long-standing

challenge, and still a key area of interest, lies in understanding the factors controlling inter–annual variability (IAV) of the various carbon fluxes (D. Baldocchi et al., 2018). The challenge presents itself from the mechanistic to the more data–driven approaches and contests the dominant role of meteorology in determining the IAV of ecosystem fluxes (Richardson et al., 2007). At the local ecosystem level, J. Wu et al. (2012) looked at the IAV of net ecosystem fluxes by fitting the parameters of a semi-empirical model at shorter timescales to capture the seasonality, but also annual variability of model parameters. The approach allows testing the role of changes in ecosystem functioning in the IAV of carbon fluxes (Richardson et al., 2007). They concluded that climate and parametric variability control IAV of ecosystem fluxes at shorter and longer timescales, respectively.

Simultaneously, Fatichi and Ivanov (2014) highlighted the role of climate when using 200 years of hourly synthetic meteorological data to force an ecohydrological model to find that the random occurrence of favorable weather conditions at certain hours of the day can be a major predictor of IAV of net primary production (NPP). This statistical relationship was corroborated by Zscheischler et al. (2016) using actual flux data from EC sites from forested areas in North America, where the 91st percentile values of hourly GPP flux, that is, peak GPP values, substantially contributed to the IAV of GPP flux. These studies highlight the correlation between the distribution tails and the IAV in EC fluxes. However, there is no robust pattern across sites nor do they challenge there is no variability in ecosystem function.

More recently, a model selection study compared an ensemble of 5600 possible semi-empirical LUE model structures to find a global best model structure (Bao, Wutzler, et al., 2022). The best LUE model was calibrated at a daily timescale per site and simulated GPP fluxes across the FLUXNET EC tower sites (Pastorello et al., 2020), considering the effect of various environmental conditions on maximum LUE through partial sensitivity functions. Though the best global model performed similarly to the best model selected for each site at the daily resolution, it failed to represent the variability of annually aggregated GPP fluxes for 74% of sites, that is, the Nash-Sutcliffe efficiency (NSE) of model performance (Nash & Sutcliffe, 1970) was below or equal to 0.5. This may be attributed to (a) the use of daily data in the study, as the model had no information on the favorable conditions that occurred in a diurnal cycle and failed to simulate the diurnal GPP peaks which had a major influence on IAV (Bao, Wutzler, et al., 2022; Fatichi & Ivanov, 2014; Zscheischler et al., 2016), (b) the assumption of invariance in ecosystem function, that is, values of model parameters remain constant for all site–years in a site, and (c) the need to explicitly consider different timescales in the cost function (Desai, 2010).

In contrast, Mengoli et al. (2022) proposed an optimality-based framework (Stocker et al., 2020; Wang et al., 2017), that is, optimality-based P-model which simulates GPP following optimality principle and differentiates between instantaneous and acclimated photosynthetic responses. This model demonstrated its capability in simulating half-hourly GPP dynamics at 10 EC sites, covering four vegetation classes for limited time periods. Whereas, the performance of this modeling framework across sites representing diverse climate-vegetation features and various temporal resolutions were not evaluated. Though this modeling framework considers the effect of temperature, vapor pressure deficit (VPD), atmospheric CO₂ concentration, solar radiation, and the fraction of absorbed photosynthetically active radiation (fAPAR), it does not explicitly consider the effect of drought stress on GPP variability at sub-daily scale. However, numerous studies (Anderegg et al., 2015; Assal et al., 2016; Kannenberg et al., 2019; Reichstein et al., 2013; C. Wu & Wang, 2022) in the past highlighted the effect of climatic extremes, such as drought on the reduction of GPP through observations and simulations, as well as ecosystems can temporarily act as a carbon sink during drought (Ciais et al., 2005). Moreover, various studies (Müller & Bahn, 2022; Orth et al., 2020; Seneviratne et al., 2012; X. Yu et al., 2022) also found that the drought can also impose a prolonged legacy effect on seasonal and annual dynamics of GPP and recovery days can vary between ecosystems (Z. Yu et al., 2017). Recently, Mengoli et al. (2023) proposed an improved version of this model by incorporating climatic aridity and calculating a scaling factor for GPP. However, in the improved model, the scaling factor could only be applied to improve the simulation of daily GPP.

The challenge to correctly reproduce IAV is also apparent on a global scale as Anav et al. (2015) found a high level of disagreement in annual GPP estimates from diverse global GPP modeling frameworks. These discrepancies highlight that site-level limitations in simulating IAV propagate to larger scales where additional mechanisms play a role, such as natural or anthropogenic disturbances and land–use land cover change (Bultan et al., 2022; McGuire et al., 2001).

As highlighted above, the major persistent drawbacks in most of the past studies were the limited implementation of either model structures or parameterization approaches and the evaluation of models for limited sites and

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timescales. To address this, here we explore ecosystem-level estimations of GPP flux to systematically investigate how various factors can be linked to describing the IAV of GPP flux, such as peak values of diurnal GPP, climatic conditions, and variables represented by model parameters, which are usually hard to measure directly and can be difficult to interpret even when various modeling approaches are adopted. We tested the impact of the constant or time–varying parameterizations and evaluated their performances in capturing GPP variability at various temporal aggregation scales, especially at the annual scale. We also tested the hypothesis that observational constraints complement and enhance theoretically grounded process formulations and that improving the model simulations at the sub-daily scale improves the prediction of IAV of GPP. Additional analysis on parameter inversion approaches and cost functions, as well as on parametric variability are treated in a companion paper (De, Brenning, et al., 2025). In this study, we aim to answer:

- 1. How well does an optimality-based model perform compared to a semi-empirical model across various temporal scales with different model parameterization approaches?
- 2. What factors influence the variability of model performance at different temporal scales?
- 3. Can the performance of an optimality-based model be improved if drought stress is included?
- 4. How much are the differences in model performance between an optimality-based and a semi-empirical model related to variations across plant-functional types (PFT) and climate-vegetation types?
- 5. Does improved simulations of peak diurnal GPP lead to improved simulations of IAV of GPP?

2. Methods and Data

In this study, we focused on parameterization of both a semi-empirical model, at daily and sub-daily scales, and an optimality-based model at a sub-daily scale using various parameterization strategies consisting of different subsets of data and cost functions (Figure 1). Thereafter, we performed forward runs of models with calibrated parameters at the temporal resolution of model parameterization data and evaluated model performances at different temporal aggregations (Figure 1). The following sections describe each methodological step in a detailed manner.

2.1. Models

2.1.1. Optimality-Based Model: P-Model of Mengoli

Stocker et al. (2020) proposed the first version of the P-model based on theories formulated by Wang et al. (2017), which unified the classic Farquhar–von Caemmerer–Berry model (Farquhar et al., 1980) with the simplified formation of big leaf LUE models (Monteith, 1972). The probable reasons behind using the "P" in the P-model are (a) "P" stands for photosynthesis, (b) classically, GPP used to be denoted by "P" (Monteith, 1972), and (c) the initial of the lead author (Prentice et al., 2014) who formulated the theories behind the model starts with "P" (B. D. Stocker, personal communication, 06 May 2024). The underlying equations of the P-model were formulated based on the optimality principle (Prentice et al., 2014) and the coordination principle (J.-L. Chen et al., 1993; Maire et al., 2012). According to the optimality principle, plants aim to optimize the cost of transpiring water to assimilate CO₂ through the stomata. In the P-model, the ratio of leaf internal and ambient CO₂ concentration ($\chi = C_i/C_a$) is calculated for which the above-described cost is minimal, and the sensitivity (ξ) of χ to VPD is predicted. The coordination principle describes the achievement of equilibrium between the maximum rate of carboxylation ($V_{c_{mx}}$) and electron transport (J_{max}) by the plants.

Mengoli et al. (2022) adapted the first version of the P-model to simulate half-hourly GPP dynamics. Here, we applied this same model at an hourly scale and called it the P_{hr} model. The major improvement in this version was defining an explicit differentiation between instantaneous (such as RuBisCo and light-limited carbon assimilation), and photosynthetic responses ($V_{c_{max}}$, J_{max} , and ξ) which acclimate over time in response to environmental conditions. One of the important aspects of this P_{hr} model is that the parameters associated with cellular biochemistry acclimate to favorable conditions during the day over a period of time or acclimation time (A_t). In this study, we considered the favorable condition as the average of three hourly input data points in the middle of the day from 11:00 (hh:mm) LT, 12:00 (hh:mm) LT, and 13:00 (hh:mm) LT. A rolling mean of the average condition from mid-day was taken over the A_t , which was used to calculate optimal values of the model parameters, as described in Mengoli et al. (2022). The value of A_t was calibrated as a parameter in our study





Figure 1. Graphical representation of the model–data–integration workflow adopted in this study. The blue box indicates the preparation of forcing and observation data at hourly and daily scales for each site, as well as defines the initial value of parameters and their range by surveying literature. Then five different model parameterization tasks were performed for the light use efficiency (LUE) model from Bao, Wutzler, et al. (2022) at hourly scale (Bao_{hr} model) and at daily scale (Bao_{dd} model), P-model from Mengoli et al. (2022) at hourly scale (P_{hr} model), and P_{hr} model with an explicit drought stress function (P_{hr}^{W} model) using the Covariance Matrix Adaptation Evolution Strategy (CMA-ES) (Hansen & Kern, 2004), which is indicated by the red box. The cost function (*f*) is a function of observed (*y*) and simulated (*x*) gross primary production. The green box denotes that the whole workflow was applied for the 198 sites from the FLUXNET2015 data set (Pastorello et al., 2020). The dotted orange box highlights the focus of this study. The parameter dynamics is explored in detail in a companion paper (De, Brenning, et al., 2025). The figure was created in BioRender. R. De (2025).

(Table 1). We chose the mid-day and rolling mean approach from Mengoli et al. (2022) as it produced the best results in their evaluations of the P_{hr} model at the half-hourly scale.

One of the known limitations of the P_{hr} model is its tendency to overestimate GPP fluxes in water–limited ecosystems, as no explicit representation of soil moisture conditions was included (Mengoli et al., 2022, 2023). In order to relax such drawbacks here we used the water availability index (*WAI*) as a proxy of soil moisture (Bao, Wutzler, et al., 2022; Boese et al., 2019; Tramontana et al., 2016). The *WAI* represents the spatial and temporal dynamics in plant available water based on a simple hydrological model where storage is controlled by precipitation and evapotranspiration. We further introduced a drought stress function that additionally scaled the GPP estimates of the P_{hr} model, and we denoted this new version as the P_{hr}^W model. We calibrated 10 parameters in the P_{hr}^W model in which nine parameters were in the hydrological model and the drought stress function (Table 1). Further details on the implementation of the P_{hr}^W model, along with the drought stress function can be found in Text S1 and S2, Figures S1 and S2 of Supporting Information S1.

2.1.2. Semi-Empirical Model: Bao Model

Vegetation stores energy from absorbed solar radiation in the form of biochemical energy through photosynthesis. The efficiency of the photosynthetic apparatus in performing this energy conversion is termed as light use efficiency (ϵ). In a LUE model, GPP is calculated as the product of instantaneous ϵ , photosynthetic photon flux density (*PPFD*), and the fraction of incident photosynthetically active radiation that is absorbed by vegetation (*fAPAR*). Instantaneous ϵ reaches its maximum, that is, ϵ_{max} , when all environmental factors are optimal for photosynthesis.



Table 1

Description, Range, Initial Values, and Units of Calibrated Model Parameters

			Initial	Lower	Upper		
fX/model name	Symbol	Definition	value	bound	bound	Units	Reference
P_{hr}^{W} , P_{hr} models	A_t	Length of acclimation time	18	1	100	Days	after Mengoli et al. (2022)
Bao _{hr} , Bao _{dd} models	$\varepsilon_{\rm max}$	Maximum light use efficiency	0.04	0	0.13	$\mu molCO_2 \cdot \mu mol \ photons^{-1}$	Skillman (2008)
fT (Bao _{hr} , Bao _{dd} models)	T_{opt}	Optimal temperature	10	5	35	°C	Bao, Wutzler, et al. (2022)
	k_T	Sensitivity to temperature changes	2	1	20	$^{\circ}C^{-1}$	Bao, Wutzler, et al. (2022)
	α_{fT}	Lag parameter for temperature effect	0.29	0	0.9	-	Bao, Wutzler, et al. (2022)
fVPD (Bao _{hr} , Bao _{dd} models)	κ	Sensitivity to VPD changes	-5×10^{-5}	-0.01	-1×10^{-5}	Pa^{-1}	Bao, Wutzler, et al. (2022)
	C_{κ}	Sensitivity to atmospheric CO ₂ concentration changes	0.4	0	10	-	Bao, Wutzler, et al. (2022)
	C_{a0}	Minimum optimal atmospheric CO ₂ concentration	380	340	390	ppm	Bao, Wutzler, et al. (2022)
	C_m	CO2 fertilization intensity indicator	2000	100	4,000	ppm	Bao, Wutzler, et al. (2022)
fL (Bao _{hr} , Bao _{dd} models)	γ	Light saturation curvature indicator	2×10^{-3}	0	0.05	$\mu mol \ photons^{-1} \cdot m^2 \cdot s$	Bao, Wutzler, et al. (2022)
<i>fCI</i> (Bao _{hr} , Bao _{dd} models)	μ	Sensitivity to cloudiness index changes	0.5	10 ⁻³	1	-	Bao, Wutzler, et al. (2022)
$ \begin{array}{c} fW \ (\mathrm{P}_{\mathrm{hr}}^{\mathrm{W}}, \mathrm{P}_{\mathrm{hr}}, \\ \mathrm{Bao}_{\mathrm{hr}}, \mathrm{Bao}_{\mathrm{dd}} \\ \mathrm{models}) \end{array} $	W _I	Optimal soil moisture	0.26	0.01	0.99	$mm \cdot mm^{-1}$	Bao, Wutzler, et al. (2022)
	k_W	Sensitivity to soil moisture changes	-11	-5	-30	-	Bao, Wutzler, et al. (2022)
	α	Lag parameter for soil moisture effect	0.98	0	1	-	Bao, Wutzler, et al. (2022)
WAI (P ^W _{hr} , P _{hr} , Bao _{hr} , Bao _{dd} models)	AWC	Available water capacity	100	1	1,000	mm	Bao, Wutzler, et al. (2022)
	θ	Rate of evapotranspiration	0.05	10^{-4}	0.1	$\mathrm{mm}\cdot\mathrm{h}^{-1}$	Bao, Wutzler, et al. (2022)
	PET_{scalar}	Multiplicative scalar for potential evapotranspiration	1.2	0	5	_	Trautmann et al. (2018)
	MR _{tair}	Snow melt rate for temperature	0.125	0	0.5	$mm\cdot \ ^{\circ}C^{-1}\cdot h^{-1}$	Trautmann et al. (2018)
	MR _{netrad}	Snow melt rate for net radiation	0.0375	0	0.125	$mm\cdot MJ^{-1}\cdot h^{-1}$	Trautmann et al. (2018)
	sn _a	Sublimation resistance	0.44	0	3	_	Bao, Wutzler, et al. (2022)

Instantaneous ε is determined as the product between ε_{max} and the partial sensitivity functions (*fX*) for the different environmental factors controlling GPP, such as air temperature (*T*), *VPD*, available soil water supply (*W*), absorbed photosynthetic photon flux ($L = PPFD \times fAPAR$), the cloudiness index (*CI*, Table A1), and atmospheric CO₂ concentration (Bao, Wutzler, et al., 2022; Horn & Schulz, 2011; Mäkelä et al., 2008).

$$GPP_{sim} = (\boldsymbol{\varepsilon}_{\max} \cdot fT \cdot fVPD \cdot fL \cdot fCI \cdot fW) \cdot PPFD \cdot fAPAR$$
(1)

$$fT = \frac{2 \cdot \exp\left(-\frac{T_f - T_{opt}}{k_T}\right)}{1 + \left(\exp\left(-\frac{T_f - T_{opt}}{k_T}\right)\right)^2}$$
(2)

$$T_f(t) = \left(1 - \alpha_{fT}\right) \cdot T(t) + \alpha_{fT} \cdot T_f(t-1)$$
(3)

$$fVPD = \exp\left(\kappa \cdot \left(\frac{C_{a0}}{CO_2}\right)^{C_{\kappa}} \cdot VPD\right) \cdot \left(1 + \frac{CO_2 - C_{a0}}{CO_2 - C_{a0} + C_m}\right)$$
(4)



$$fL = \frac{1}{\gamma(PPFD \cdot fAPAR) + 1} \tag{5}$$

$$fCI = CI^{\mu} \tag{6}$$

$$fW = \frac{1}{1 + \exp\left(\boldsymbol{k}_{W}\left(\boldsymbol{W}_{f} - \boldsymbol{W}_{I}\right)\right)} \tag{7}$$

$$W_{f_t} = (1 - \alpha) \cdot W_t + \alpha \cdot W_{f_{t-1}}$$
(8)

In this study, we used the LUE model of Bao, Wutzler, et al. (2022), Bao et al. (2023) since it emerged as a robust representation from the systematic comparison across the large diversity of LUE formulations in the literature. The model selection followed a Bayesian approach that leveraged on the evaluation of modeling performance across FLUXNET EC sites (Pastorello et al., 2020) when forced and calibrated with daily data for each site. We denoted this model as the Baohr model when we parameterized at hourly scale, and as the Baodd model when we parameterized at daily scale. The model is described in Equations 1–8, where fT, fVPD, fW, fL, and fCI are partial sensitivity functions for T, VPD, W, L, and CI, respectively. In this case, W and fW were calculated similar to the implementation in P_{hr}^{W} model, that is, with a simple hydrological model (Text S1 and Figure S1 of Supporting Information S1) and drought stress function (Equations 7 and 8 are same as Equations S1 and S2 of Supporting Information \$1), respectively. Bold terms in the Equations 1-8 are model parameters, and their initial values, units, and ranges are described in Table 1. The physical ranges for most of the parameters were based on Bao, Wutzler, et al. (2022), Bao et al. (2023), and Trautmann et al. (2018). The fVPD term, viz. Equation 4, also accounts for atmospheric CO₂ concentration. The partial sensitivity functions range from zero to one (except the 2nd part of Equation 4 which can be greater than one), where a value of zero completely diminishes, and of one completely favors GPP. In this study, we changed the denominator of Equation 2 in comparison to the original exponential function $\exp\left(-\frac{T_f - T_{opt}}{k_T}\right)^2$ of Bao, Wutzler, et al. (2022), Bao et al. (2023), as the revised version produced a more realistic range of fT (Figure S3 of Supporting Information S1). Sensitivity functions fT and fWalso consider a lag effect of T and W. The lag effect of temperature was considered for Temperate, Boreal, and Polar regions where the first letter of the Köppen–Geiger (KG) climate class is "C," "D," "E," and that of soil water supply was considered for arid regions where the first letter of the KG climate class is "B" (Beck et al., 2018; Rubel et al., 2017).

2.2. Data Used

We selected 198 eddy covariance sites, for which the required forcing and observation or derived data for model parameterization were available from the FLUXNET2015 data set (Pastorello et al., 2020; FLUXNET. org, 2024a). A list of these sites is provided in Table S8 of Supporting Information S1 and their spatial distributions are plotted in Figure S4 of Supporting Information S1. The variables which were used to force, and parameterize models as well as data processing steps such as gap–filling, and quality control are described in detail in Table A1, Appendix B and Appendix C. We prepared these data in both hourly and daily resolutions.

In our study, a total of 13 different PFTs (as defined in FLUXNET.org, 2024b) were represented: croplands (CRO; 19 sites), deciduous broadleaf forests (DBF; 25 sites), deciduous needle leaf forest (DNF; one site), evergreen broadleaf forests (EBF; 13 sites), evergreen needle leaf forests (ENF; 47 sites), grasslands (GRA; 35 sites), mixed forests (MF; nine sites), closed shrublands (CSH; three sites), open shrublands (OSH; 13 sites), savannas (SAV; six sites), permanent wetlands (WET; 20 sites), woody savannas (WSA; six sites), and land cover under snow for most of the year (SNO; one site). The major KG climate classes (Rubel et al., 2017; Beck et al., 2018; FLUXNET.org, 2024c) are represented by 12 tropical sites, 18 arid sites, 87 temperate sites, 71 boreal sites, and 10 polar sites. We also classified sites into nine climate–vegetation types, similar to Bao, Wutzler, et al. (2022), in which seven sites are tropical forests (TropicalF), five sites are tropical grassland (TropicalG), six sites are arid forest (AridF), 12 sites are arid grassland (AridG), 51 sites are temperate forest (TemperateF), 36 sites are temperate grassland (TemperateG), 52 sites are boreal forest (BorealF), 19 sites are boreal grassland (BorealG), and 10 sites have polar vegetation.



Table 2

Description of Models and Tasks Accomplished With Each Specific Model

		Parameterization strategies				
Models	Description	Per site-year	Per site using Cost ^{IAV}	Per site	Per PFT	Global
P _{hr}	P-model of Mengoli et al. (2022) parameterized using hourly data	a, d	a, d	a, d	a, d	a, d
$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	P-model of Mengoli et al. (2022) with an additional constraint on drought stress and parameterized using hourly data	a, b, d, e	a, b, d, e	a, b, d, e	a, b, d, e	a, b, d, e
Bao _{hr}	LUE model of Bao, Wutzler, et al. (2022) parameterized using hourly data	a, c, d, e	a, c, d, e	a, c, d, e	a, c, d, e	a, c, d, e
Bao _{dd}	LUE model of Bao, Wutzler, et al. (2022) parameterized using daily data	a, c, d	a, c, d	a, c, d	a, c, d	a, c, d

Note. The tasks are described in the footnote of the table. ^aEvaluation of model performance across timescale with different model types, parameterization strategies, and cost functions. ^bEvaluation of an optimality-based model with an explicit drought stress function. ^cEvaluation of a semi-empirical model with different temporal resolutions of data used for model parameterization. ^dFactors behind variability of model performance across timescales. ^eVariability of annual model performance with model performance in simulating diurnal gross primary productivity (GPP) peaks.

2.3. Model Parameterization

We primarily defined four different parameterization strategies consisting of various subsets of data to calibrate the model parameters controlling hourly GPP dynamics. These parameterization strategies were used to determine a vector of calibrated parameter values (a) for each site–year, (b) for each site, (c) for each PFT, and (d) for all sites at once (global parameterization). We also performed another parameterization per site using a modified cost function which used an additional constraint on the IAV of GPP ($Cost^{IAV}$). We chose various parameterization strategies ranging from more flexible to rigid to evaluate how good or bad the model performs with changing parameterization. We parameterized and forced the Bao_{hr} model, and the Bao_{dd} model using hourly and daily data, respectively to perform a comparative analysis (Table 2). The P^W_{hr} model and the P_{hr} model were only parameterized and forced using hourly data (Table 2).

We used Python (Python Core Team, 2021) implementation (pycma v3.3.0.1) of the Covariance Matrix Adaptation Evolution Strategy (CMA-ES) (Hansen & Kern, 2004; Hansen et al., 2019) as our global search algorithm to find the values of model parameters for which cost function reached its minimum. This is a derivative-free, evolutionary algorithm, which is designed to find global minima in a rugged parameter space.

$$Cost_{i_P} = (1 - GPP_{NNSE_i}) + (1 - ET_{NNSE_i})$$

$$(9)$$

$$Cost_{i_{Roo}} = (1 - GPP_{NNSE_i}) + (1 - ET_{NNSE_i}) + Cost_{ideal} + Cost_{non_ideal}$$
(10)

A robust cost function is a necessity for the numerical optimizer to find the global minimum. The cost functions for P_{hr}^W , P_{hr} models ($Cost_{i_p}$) and the Bao_{hr}, Bao_{dd} models ($Cost_{i_{Bao}}$) were calculated as Equations 9 and 10, respectively, in case of per site–year and per–site parameterization. Here, *i* is either a site or site–year based on parameterization type. For PFT-specific model parameterization, the cost functions were $\sum_{i=1}^{N_{PFT}} Cost_{i_{pao}}$ for P_{hr}^W , P_{hr} models and Bao_{hr}, Bao_{dd} models, respectively. *i* denotes a site and N_{PFT} denotes the total number of sites in a specific PFT. In the case of global model parameterization, the cost functions were $\sum_{i=1}^{N} Cost_{i_p}$ and $\sum_{i=1}^{N} Cost_{i_{Bao}}$ for the P_{hr}^W model and Bao_{hr}, Bao_{dd} models, respectively. *i* denotes a site and *N* denotes the total number of sites used in this study.

$$NNSE_i = \frac{1}{2 - NSE_i} \tag{11}$$



$$NSE_{i} = 1 - \frac{\sum_{t=1}^{N_{t,i}} (\sigma_{weight_{t,i}} \cdot (EC_{t,i} - sim_{t,i}))^{2}}{\sum_{t=1}^{N_{t,i}} (\sigma_{weight_{t,i}} \cdot (EC_{t,i} - \overline{EC_{t,i}}))^{2}}$$
(12)

$$\sigma_{weight_{t,i}} = 1 - \frac{\sigma_{t,i} - \min(\sigma_i)}{\max(\sigma_i) - \min(\sigma_i)}$$
(13)

GPP_{NNSE}, and ET_{NNSE}, were calculated (Equation 11) as a weighted normalized NSE, viz. NNSE (Hundecha & Merz, 2012) between the time series of good quality data points (see Appendix B for the selection criteria) of EC derived and simulated GPP and ET, respectively. The GPP and ET derived from EC measurements are denoted as GPP_{EC} and ET_{LE} , respectively. The simulated GPP and ET are denoted as GPP_{sim} and ET_{sim} (see Figure S1 of Supporting Information S1 for calculation of ET_{sim}), respectively. We considered ET as well in our cost function to better constrain the parameters of the simple hydrological model used in this study. The NNSE values (Nash & Sutcliffe, 1970) are between zero and one, where one is the best, and zero is the worst agreement between observed and simulated data. Here, we used these normalized values so that minimizing (1 - NNSE) always results in better model performance in comparison to using (1 - NSE), where NSE can have values between $-\infty$ (worst agreement) and one (best agreement). In Equation 12, $N_{t,i}$ is the total number of good quality data points from each timestep t for a site-year or site i. $\sigma_{t,i}$ in Equation 13 is random uncertainty (standard deviation of fluxes in a sliding window of ± 5 days and ± 1 hr of the time-of-day of the current timestamp) of NEE or ET from each timestep t for a site-year or site i (Table A1). Similarly, σ_i in Equation 13 is random uncertainty of full NEE and ET time series for a site-year or site *i*. The normalized random uncertainty in Equation 13 was used in the cost function to allocate higher and lower weight to EC-derived values with lower and higher uncertainties, respectively.

$$Cost_{ideal} = ((1 - \max(fT_r)) + (1 - \max(fVPD_r)) + (1 - \max(fW_r)) + (1 - \max(fL_r))) \cdot 10^3$$
(14)

$$Cost_{non_ideal} = \sum_{r} \left(\left(fT_r - \theta_{fT} \right) \left(T < 0^{\circ} C \& fT_r > \theta_{fT} \right) \right)$$

$$+ \sum_{r} \left(\left(fVPD_r - \theta_{fVPD} \right) \left(VPD > 2000Pa \& fVPD_r > \theta_{fVPD} \right) \right)$$

$$+ \sum_{r} \left(\left(fW_r - \theta_{fW} \right) \left(W < 0.01 \& fW_r > \theta_{fW} \right)$$

$$(15)$$

The *Cost_{ideal}* and *Cost_{non_ideal}* were introduced as regularizers in *Cost_{ibao}* to avoid over-fitting of the sensitivity functions (Bao et al., 2023; Bao, Wutzler, et al., 2022). These cost function components ensure that values of partial sensitivity functions were not penalized and favored under ideal and non-ideal conditions, respectively. The ideal and non-ideal conditions were determined by certain constant thresholds for all sites. Equation 14 ensured that the partial sensitivity functions, *fT* (Equation 2), only left part of the *fVPD* (Equation 4), *fW* (Equation 7) and *fL* (Equation 5) approaches one, when certain ideal environmental conditions (*PPFD* \in [0 to 600 µmol photons $\cdot m^{-2} \cdot s^{-1}$], *fAPAR* \in [0 to 1], *T* \in [-5 to 40 °C], *VPD* \in [0–4,500 Pa], *W* \in [0 to 1]) occur (these ranges are denoted by subscript *r*), so that the ε_{max} in Equation 1 reaches its maximum potential. The factor 10³ in Equation 14 was included to match the ranges of all other components in the cost function for the Bao_{hr}, Bao_{dd} models (*Cost_{ibao}*) so that all the components had equal weight. Equation 7), were greater than a certain threshold ($\theta_{fT} = 0.2$, $\theta_{fVPD} = 0.9$, $\theta_{fW} = 0.2$) under non-ideal conditions (*T* < 0 °C, *VPD* > 2,000 Pa, *W* < 0.01) for photosynthesis.

$$Cost_{i_{p}}^{IAV} = (1 - GPP_{NNSE_{i}}) + (1 - GPP_{NNSE_{i}}^{y}) + (1 - ET_{NNSE_{i}})$$
(16)

$$Cost_{i_{Ren}}^{IAV} = (1 - GPP_{NNSE_i}) + (1 - GPP_{NNSE_i}^{y}) + (1 - ET_{NNSE_i}) + Cost_{ideal} + Cost_{non_ideal}$$
(17)



$$GPP_{NNSE_i}^y = \frac{1}{2 - GPP_{NSE_i}^y}$$
(18)

$$GPP_{NSE_{i}}^{y} = 1 - \frac{\sum_{t=1}^{N_{t,i}} (\sigma_{weight_{t,i}}^{y} \cdot (EC_{t,i}^{y} - sim_{t,i}^{y}))^{2}}{\sum_{t=1}^{N_{t,i}} (\sigma_{weight_{t,i}}^{y} \cdot (EC_{t,i}^{y} - \overline{EC_{t,i}^{y}}))^{2}}$$
(19)

$$\sigma_{weight_{t,i}}^{y} = 1 - \frac{\sigma_{t,i}^{y} - \min(\sigma_{i}^{y})}{\max(\sigma_{i}^{y}) - \min(\sigma_{i}^{y})}$$
(20)

$$EC_{t,i}^{y} = \sum_{t=1}^{t} EC_{t,y,i}; \ sim_{t,i}^{y} = \sum_{t=1}^{t} sim_{t,y,i}; \ \sigma_{t,i}^{y} = \sum_{t=1}^{t} \sigma_{t,y,i}$$
(21)

In the case of per–site–year parameterization using cost functions in Equations 9 and 10, we fitted the model so that the annual average of GPP can be captured well for each site–year. Whereas, in the case of per–site parameterization using cost functions in Equations 9 and 10, the model was parameterized for each site. We performed another experiment as a balance between these two experiments using the $Cost^{IAV}$, which is similar to Desai (2010) to put an additional constraint on IAV, and parameterized P_{hr}^W , P_{hr} , Bao_{hr}, and Bao_{dd} models for each of the EC sites. The cost functions, $Cost^{IAV}_{ip}$ for P_{hr}^W , P_{hr} models (Equation 16) and $Cost^{IAV}_{igao}$ for Bao_{hr}, and Bao_{dd} models (Equation 17) now include an additional term (1 – GPP_{NNSE_i} , ^y) to constrain the annual cumulative sum of GPP flux from each site *i*. $EC^y_{i,i}$, $sim^y_{i,i}$, and $\sigma^y_{i,j}$ (Equation 21) are cumulative sums of GPP_{EC} , GPP_{sim} , and σ_{NEE} from start of each year *y* to timestep *t* for each site *i*, respectively.

2.4. Simulating and Evaluating GPP Estimates

2.4.1. Forward Runs

In the case of the site–year parameterization, we performed a forward run for each site–year using the respective set of calibrated parameter values and forcing data for that year. Afterward, we concatenated GPP_{sim} from all the years for a given site to assess model performance. For per–site parameterization using $Cost^{IAV}$, and per–site parameterization, we used site-specific values of calibrated parameters to perform site-level model evaluation. We also applied calibrated model parameterization, a single set of calibrated parameter values was used to simulate GPP for each site.

2.4.2. Model Performances Metric

We performed forward runs at an hourly scale and averaged the hourly simulations to daily, weekly, monthly, and annual temporal frequencies to calculate model performance measures at different temporal aggregations. Model performance was only evaluated for temporal aggregations from daily to annual for the Bao_{dd} model. We applied a data screening procedure (Appendix C) before calculating model performance measures.

NSE(y,
$$\hat{y}$$
) = 1 - $\frac{\sum_{t=1}^{N} (y_t - \hat{y}_t)^2}{\sum_{t=1}^{N} (y_t - \bar{y})^2}$ (22)

We evaluated how well a model can simulate the IAV of GPP based on how well a model simulated the annual average GPP for a site. In this study, we performed most of our analysis using NSE (Nash & Sutcliffe, 1970) and normalized NSE, viz. NNSE (which is $\frac{1}{2-NSE}$) as NSE indicates the degree to which scatter between observed and simulated data fits to the 1:1 line. We used Equation 22 to calculate NSE between a reference (*y*) and a predicted (\hat{y}) variable, where *N* is the total number of time-steps, and y_t , \hat{y}_t are the values of reference and predicted variables, at timestep *t*. In addition, we calculated the square of the Pearson correlation coefficient (\mathbb{R}^2) (Kirch, 2008) which explains whether the dispersion of observed and simulated data matches and in the case of an unbiased model, values of NSE will be closer to values of \mathbb{R}^2 . Whereas, if a model is systematically biased, it will result in higher \mathbb{R}^2 values, but bad NSE values (Krause et al., 2005). We also calculated Root Mean Squared Error



(Chai & Draxler, 2014) to quantify how closely the mean of simulated data matches with the mean of the observed data.

$$NSE = 2 \cdot \alpha_{NSE} \cdot r - \alpha_{NSE}^2 - \beta_n^2 \tag{23}$$

$$\alpha_{NSE} = \frac{\sigma_{sim}}{\sigma_{EC}} \tag{24}$$

$$\beta_n = \frac{\mu_{sim} - \mu_{EC}}{\sigma_{EC}} \tag{25}$$

Moreover, using Equations 23–25, we decomposed NSE values to linear correlation (*r*), relative variability (α_{NSE}), and bias (β_n) in some cases to investigate which of these were improved or diminished between different model parameterization strategies (Gupta et al., 2009). In Equations 24 and 25, σ_{sim} and σ_{EC} are standard deviations of GPP_{sim} and GPP_{EC} , respectively, μ_{sim} and μ_{EC} are mean GPP_{sim} and GPP_{EC} , respectively. We calculated these metrics using the Python (Python Core Team, 2021) package Permetrics v1.5.0 (Van Thieu, 2023; Van Thieu & Mirjalili, 2023), and the definition of each of the model performance metrics can be found in the package documentation.

2.4.3. Estimating Uncertainties in Annual GPP

GPP is not directly measured using the EC technique (Foken et al., 2011; Montgomery, 1948; Swinbank, 1951). Rather, GPP is derived from NEE using a flux partitioning method (Lasslop et al., 2010; Reichstein et al., 2005). The measurement of NEE is affected by low turbulence conditions when the EC system misses the carbon flux due to advection transport (Aubinet et al., 2010). The low turbulence conditions were identified by friction velocity threshold (u_*), and NEE measurements during these conditions were discarded and gap-filled (Papale et al., 2006). 200 values of u_* thresholds were estimated using a bootstrap approach for each site–year, and finally 40 different values (chosen from percentile 1.25 to percentile 98.75 with a step of 2.5) of u_* was used to produce 40 different realizations of NEE data. These u_* thresholds were either estimated per year using variable u_* threshold (VUT) method where the u_* threshold of a year depends on data from a given year and its neighboring years or estimated using data from all years, and remains constant across years, which is known as constant u_* threshold (CUT) method (Pastorello et al., 2020). Thereafter, seven different representative NEE data were produced as percentiles (XX) 5, 16, 25, 50, 75, 84, and 95 of the 40 different NEE values for each VUT and CUT method. These seven NEE data were further partitioned to estimate GPP using daytime (DT) and night-time (NT) partitioning methods (Lasslop et al., 2010; Reichstein et al., 2005).

Here, we used GPP_NT_VUT_USTAR50 from the FLUXNET2015 data set (Pastorello et al., 2020) as our reference GPP or GPP_{EC} for model calibration and evaluation (see Section 2.2 and Appendix A). These GPP values were derived using night-time partitioning from NEE data (NEE_VUT_USTAR50) where the VUT method was considered, and the 50 percentile value of u_{*} thresholds was applied. However, the GPP data disseminated in the FLUXNET2015 data set has a systematic uncertainty on the choice of NEE data (depending on percentile XX), and partitioning methods. This uncertainty can be large when the fluxes are aggregated to annual scale. We annually aggregated four other GPP variables (GPP_NT_VUT_05, GPP_NT_VUT_95, GPP_DT_VUT_05, GPP_DT_VUT_95) derived using either daytime or night-time partitioning method, and derived from 5 (NEE_VUT_05) and 95 (NEE_VUT_95) percentile NEE values, and screened for good site-years with their respective quality control flags (see Appendix C). We quantified the uncertainty range of annual GPP_{EC} for a site-year as the minimum of annual GPP_NT_VUT_05 and GPP_DT_VUT_05 to the maximum of annual GPP_DT_VUT_95 and GPP_NT_VUT_95. Thereafter, we also quantified the fraction of site years in a site for which annual GPP_{sim} estimated from different experiments were within the uncertainty range of GPP_{EC} . The values of GPP_{EC} were also affected by random uncertainties in NEE flux. We used hourly random uncertainties in NEE flux to weigh the contribution of a GPP value from a given hour in the cost function (Section 2.3). However, the random uncertainty at annual scale was not considered as they greatly reduce with aggregation (Hollinger & Richardson, 2005; Tramontana et al., 2016).



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2.4.4. Factors Influencing Model Performance

We selected potential factors that can affect model performance at different temporal resolutions. These factors can be of two types. There were factors which we determined based on our experiment design, which included model types (P_{hr}^{W} model, P_{hr} model, Bao_{hr} model, and Bao_{dd} model), parameterization strategies (per site–year, per site using *Cost^{IAV}*, per site, per PFT, and global parameterization), number of years with good quality data (Appendix C) in a site. Whereas, other factors represent site-specific characteristics, including PFT, KG climate class, and climate–vegetation types.

First, we conducted Levene's test (Levene, 1960) to determine if the assumption of homoscedasticity was fulfilled across groups in the controlling factors. Then, we performed an N-way Analysis of Variance (ANOVA) (Kaufmann & Schering, 2014) with the potential controlling factors to determine which of them played a major role in determining model performance at hourly and annual temporal scales. For analysis at an hourly scale, the Bao_{dd} model was not included as this model produced simulations at a daily scale. We performed two N-way ANOVA analyses once including the performance of the P_{hr} model, and then excluding the performance of the P_{hr} model. The Levene's test and N-way ANOVA analyses were implemented using SciPy v1.11.3 (Virtanen et al., 2020) and statsmodels v0.14.0 (Seabold & Perktold, 2010), respectively.

2.4.5. Evaluating GPP Estimates in Water-Limited Ecosystems

We investigated to determine whether explicit accounting of the drought stress function in the P_{hr}^W model had improved its performance at arid sites. For this purpose, we chose the aridity index (AI) to determine which sites were arid or semi-arid, as this index provided a numerical representation of moisture availability (Zomer et al., 2022) at a location. The AI values were calculated by dividing the average precipitation (*P*) per hour by the average potential evapotranspiration (*PET*) per hour for the whole observation period at a site.

We drew examples from a few site-specific results to highlight different aspects of the behaviour of P_{hr}^W and P_{hr} models for ecosystems with contrasting soil moisture controls on GPP and with a larger availability of good– quality measurements. For this purpose, we chose a water–limited semi-arid site (annual average precipitation of 318 mm) in central Australia (Alice Springs, AU-ASM). This site also features a complex mixture of Mulga woodland and savanna (Cleverly et al., 2013; Pastorello et al., 2020). In contrast, we also highlighted the behaviours of P_{hr}^W and P_{hr} models in an irrigated cropland (Mead - irrigated continuous maize site, US-Ne1) in the mid-western U.S.A (Amos et al., 2005; Pastorello et al., 2020).

2.4.6. Effect of Temporal Resolution of Data on Model Performance

We parameterized the LUE model of Bao, Wutzler, et al. (2022) with hourly and daily data for Bao_{hr} model and Bao_{dd} model, respectively. We performed a comparison between these two versions of the model to highlight whether the resolution of data used for model parameterization can substantially affect the prediction of the annual average or IAV of GPP fluxes. Here, we also drew a site-specific example from an energy–limited deciduous forest in central Germany (Hainich, DE-Hai) as this site had a very long observation period (Knohl et al., 2003).

2.4.7. Evaluating Modeling Experiments of Various Complexities

We formulated our experiments using models and parameterization strategies consisting of varying numbers of model parameters to be calibrated. The number of parameters calibrated for a detailed parameterization strategy, such as per site–year parameterization was substantially higher than a generic parameterization strategy, such as global parameterization. We used Akaike's Information Criterion (AIC) to investigate whether a complex modeling experiment with a higher number of parameters can better simulate GPP (Burnham & Anderson, 2004).

$$AIC = n \log\left(\frac{\sum (EC_i - sim_i)^2}{n}\right) + 2K$$
(26)

$$AIC_{c} = n \log\left(\frac{\sum (EC_{i} - sim_{i})^{2}}{n}\right) + 2K + \frac{2K(K+1)}{n - K - 1}$$
(27)

õ

governed by the applicable Crea



Following recommendations of Burnham and Anderson (2004), we used Equation 26 to calculate AIC when n/K > 40, where *n* is the total number of observations and *K* is the total number of parameters. Otherwise, we used a corrected version of AIC (AIC_c, Equation 27). Though the values of AIC or AIC_c can be in any range, the lowest value of AIC or AIC_c determines the preferred modeling experiments. *EC_i* and *sim_i* are *i*th observations of EC-derived GPP and simulated GPP, respectively in Equations 26 and 27. We considered *GPP_{sim}* from all the four variations of models, that is, P_{hr}^{W} model, P_{hr} model, Bao_{hr} model, and Bao_{dd} model for calculation of AIC or AIC_c. We calculated AIC at hourly and daily aggregations by concatenating good quality (Appendix C) hourly or daily data, and daily averages *GPP_{EC}* and *GPP_{sim}* from all the days from all sites. Similarly, we used monthly and annual aggregation, as *n* was usually smaller than *K* in these cases. The value of *K* was the total number of model parameters calibrated for all the site–years, for all the sites, for all the PFT, and for a specific model in case of per site–year parameterization, per site parameterization, respectively.

2.4.8. Difference in Model Performances Between PFTs

We studied the distribution of model performances of the P_{hr}^W model and the Bao_{hr} model for different parameterization strategies across PFTs and climate–vegetation types (Section 2.2). We performed non-parametric statistical significance testing using a two-sample Kolmogorov-Smirnov (K-S) test (Hodges, 1958) between a given pair of parameterization strategies for a given model and PFT to test if model performance obtained from a parameterization strategy is significantly different from another for a PFT and a model. Similarly, we performed the K-S test between model performances of the P_{hr}^W model and the Bao_{hr} model for a given parameterization strategy and a PFT to test if model performances were significantly different between two models. We performed the K-S test using SciPy v1.14.1 (Virtanen et al., 2020).

2.4.9. Simulating GPP Peaks

We assessed model performance in predicting peak GPP_{EC} . We defined peak GPP_{EC} and peak GPP_{sim} as the 90th percentiles of hourly GPP_{EC} ($P90_{GPP_{EC}}$) and GPP_{sim} ($P90_{GPP_{sim}}$), respectively, following the concept of good hours by Zscheischler et al. (2016) and Fatichi and Ivanov (2014). We calculated $P90_{GPP_{EC}}$ and $P90_{GPP_{sim}}$ for each site–year considering only good quality hourly data (Appendix B). We compared the ratios of peak GPP_{sim} from P_{hr}^{W} model and Bao_{hr} model to GPP_{EC} for each parameterization strategy in order to identify possible biases.

Δ

$$\Delta NNSE_{P90} = NNSE_{P90}^{j1} - NNSE_{P90}^{j2}$$
(28)

$$\Delta NNSE_{\nu} = NNSE_{\nu}^{j1} - NNSE_{\nu}^{j2}$$
⁽²⁹⁾

We furthermore investigated whether improving the simulation of peaks of GPP_{EC} improved the simulation of IAV of GPP. We calculated NNSE between $P90_{GPP_{EC}}$ and $P90_{GPP_{sim}}$ ($NNSE_{P90}^{j}$) from all the site-years in a site considering only good site-years and only for sites with more than 3 years of good quality data (Appendix C) for a parameterization strategy *j*. Similarly, we calculated NNSE between the annual average of GPP_{EC} and GPP_{sim} ($NNSE_{y}^{j}$) for sites with more than 3 years of good quality data (Appendix C) for a parameterization strategy *j*. Then, differences between $NNSE_{p90}^{j1}$ and $NNSE_{y}^{j2}$ were calculated for a pair of parameterization strategies where *j*1 and *j*2 are two different parameterization experiments, for both P_{hr}^{W} model and Bao_{hr} model (Equations 28 and 29). Correlation between $\Delta NNSE_{P90}$ and $\Delta NNSE_{y}$ were then investigated to study whether a certain parameterization strategy for a given model better captured the GPP_{EC} peaks, and thus contributed to higher annual model performance.

3. Results

3.1. Overall Model Performance

All four models, that is, P_{hr}^W , P_{hr} , Bao_{hr} , and Bao_{dd} models performed significantly better at the hourly scale than the annual scale (Figure 2). The use of an additional constraint on IAV, that is, *Cost^{IAV}* did not contribute to better





Figure 2. Distributions of model performance measure (normalized Nash-Sutcliffe efficiency, viz. NNSE) at hourly/daily scale (first row) and at annual timescale (second row) from P-model of Mengoli et al. (2022) with drought stress, parameterized at hourly scale (P_{hr}^W), P-model of Mengoli et al. (2022) without drought stress, parameterized at hourly scale (P_{hr}^W), P-model of Mengoli et al. (2022) without drought stress, parameterized at hourly scale (P_{hr}), the global best model of Bao, Wutzler et al. (2022) parameterized at hourly scale (Bao_{hr}), and global best model of Bao, Wutzler et al. (2022) parameterized at hourly scale (Bao_{hr}), and global best model of Bao, Wutzler, et al. (2022) parameterized at daily scale (Bao_{dd}). For the Bao_{dd} model, subplot (d) shows model performance at daily scale as this model was parameterized at daily scale. *Cost^{IAV}* denotes the usage of an additional constraint on annual gross primary production flux during per–site parameterization. The dotted vertical lines represent the median model performances, which are summarized in Table D1. The numbers in parentheses beside the model name on top of each of the sub-figures represent the total number of sites. The model performance at an annual scale was calculated for fewer sites as some sites have a very low measurement period (Appendix C).

model performance across sites at an annual scale and performed closer to parameterization per site and poorer than site–year parameterization (Figure 2). The median model performance was highest for the model parameterization per site–year among all model parameterization strategies (per site–year, per site using *Cost^{IAV}*, per site, per PFT, and global parameterization) for all four models (Table D1). Model parameterization per site–year also produced the best model performance at all temporal aggregation levels including annual aggregation (Figures 2, S5, and Table S2 of Supporting Information S1). P^W_{hr} model performed substantially better for the majority of the sites compared to P_{hr} model at all temporal aggregation levels as it explicitly considered site-specific water availability (Figures 2, S5, and Table S2 of Supporting Information S1). Comparison of model performances at different temporal aggregations also revealed that Bao_{hr} and Bao_{dd} models performed slightly better than the P^W_{hr} model across all timescales (hourly, daily, weekly, monthly, and annual), as the Bao_{hr} and Bao_{dd} models were more flexible than the P^W_{hr} model and captured ecosystem response with a broad range of parameters (Figure 2, Table D1, Figure S5 and Table S2 of Supporting Information S1). For example, the median NNSE(s) at the hourly resolution were 0.827 and 0.855 for the P^W_{hr} model and the Bao_{hr} model, respectively. Conversely, at the annual resolution, the median NNSE(s) were 0.543 and 0.628 for the P^W_{hr} model and Bao_{hr} model, respectively.

The annual average of GPP_{sim} was within the uncertainty range of GPP_{EC} (as defined in Section 2.4.3) for most site-years in most of the sites when P_{hr}^W (mean fraction of site-years per site: 0.87), Bao_{hr} (mean fraction of siteyears per site: 0.93), and Bao_{dd} (mean fraction of site-years per site: 0.94) models were parameterized per siteyear (Figure 3 and Table E1). P_{hr} model overestimated annual GPP_{sim} beyond the uncertainty range of GPP_{EC} in most cases for all parameterization methods, where the mean fraction of site-years per site within the uncertainty range of GPP_{EC} was ≤ 0.13 (Figure 3 and Table E1). Mean fractions of the site-years per site for which GPP_{sim} was within the uncertainty range were also similar for site-specific parameterization compared to site-year specific parameterization for P_{hr}^W , Bao_{hr}, and Bao_{dd} models (Table E1). However, for PFT-specific





Figure 3. Distributions of fractions of site–years of a site for which annual average of simulated gross primary production (GPP_{sim}) was within the uncertainty range of eddy-covariance derived GPP (GPP_{EC}) . The GPP_{sim} were estimated using (a) P-model of Mengoli et al. (2022) with drought stress, parameterized at hourly scale (P_{hr}^{hr} model), (b) P-model of Mengoli et al. (2022) without drought stress, parameterized at hourly scale (P_{hr} model), (c) global best model of Bao, Wutzler, et al. (2022) parameterized at hourly scale (B_{hr} model), (c) global best model of Bao, Wutzler, et al. (2022) parameterized at hourly scale (Bao_{hr} model), and (d) global best model of Bao, Wutzler, et al. (2022) parameterized at daily scale (Bao_{dd} model). $Cost^{IAV}$ in the legend denotes the usage of an additional constraint on annual gross primary production flux during per–site parameterization. The dotted vertical lines represent the mean fractions of site–years of a site for which the annual average of GPP_{sim} was within the uncertainty range of GPP_{EC} . The mean values are summarized in Table E1. The numbers in parentheses beside the model name on top of each of the sub-figures represent the total number of sites. The model performance at an annual scale was calculated for fewer sites as some sites have a very low measurement period (Appendix C).

parameterization and global model parameterization, the annual average of GPP_{sim} was unreliable and went beyond the uncertainty range of GPP_{EC} for most cases (Figure 3).

3.2. Factors Influencing Variability in Model Performance

We summarized the percentage contributions of factors which influenced model performance at hourly and annual scales and found that most of the variability in model performance came from how we designed our modeling experiments (Figure 4). Model types was a crucial factor when the P_{hr} model was included in N-way ANOVA analysis, as this model had comparatively poor performance at both hourly and annual scales and resulted in greater variability in the NNSE values (70.1% and 63.5% contribution to the sum of squares of the regression, viz. SSR in hourly and annual scale, respectively). We then excluded the P_{hr} model from further analysis to uncover the other factors behind the model performance and found that for the hourly scale, the model performance varied the most across the groups of KG classes (33.9% contribution to the SSR), followed by parameterization type (31.0% to the SSR) and climate–vegetation type (25.4% contribution to the SSR) (Figure 4). However, at an annual scale, the parameterization strategy strongly affected (62.3% contribution to the SSR) the model performance, as per–site–year parameterization usually better simulated the annual GPP_{obs} compared to other parameterization strategies. The number of good years (Appendix C) used for calculating annual NNSE also exerted a small influence (4.9% contribution to the SSR) on the annual model performance. In general, there were only slight performance differences between models when the P_{hr} model was not considered, and model parameterization played a bigger role in the variability of model performance.





Figure 4. Percentage contributions of factors influencing variability in model performance (normalized Nash-Sutcliffe efficiency, viz. NNSE) in the sum of squares in N-way Analysis of Variance (ANOVA). The percentage contributions show the influence of various factors on hourly and annual model performance when the P-model (Mengoli et al., 2022) without any explicit drought stress function, parameterized at hourly scale (P_{hr} model) was considered in the analysis, as well as on hourly and annual model performance excluding the P_{hr} model. The sum of squares of residual was removed before plotting the percentage contributions of the factors and only the explained variance is shown.

3.3. Effect of Drought Stress on Model Performance

The performance of the P_{hr}^W model to predict the annual average of GPP_{EC} from each of the site-years substantially improved in comparison to P_{hr} model (from an NSE of -1.39 to 0.92) after explicit consideration of soil water supply in the model (Figure 5). Most of this improvement came from the better prediction of the annual average of GPP_{EC} at the arid and semi-arid sites (with AI values lower than 0.5). For the semi-arid site (AU-ASM) the predicting performance of the P_{hr}^W model for all the parameterization strategies largely benefited from the explicit inclusion of soil water supply constraints (Figure 6). The systematic bias in model simulations was also improved after the inclusion of a drought stress constraint, as well as the modeling bias also improved from a generalized to a detailed parameterization strategy. Although the coupling of a simple hydrological model which calculated water–availability based on precipitation and evapotranspiration and inclusion of drought stress function generally improved the P_{hr}^W model for most of the site–years at an arid site, the model failed to capture the GPP_{EC} (Text S3, Figure S6 of Supporting Information S1) at an irrigated cropland site (US-Ne1), as the simple hydrological model which we used to calculate water–availability lacked representation of human management.

3.4. Effect of Temporal Resolution of the Data on Model Performance

The use of hourly data to constrain Bao_{hr} model parameters and aggregating hourly values of GPP_{sim} to annual scale did not have a significant effect on the Bao_{hr} model performance in comparison to parameterization of the same model with daily data, that is, Bao_{dd} model, in simulating the annual average of GPP_{EC} (Figure 7) for each site–year. The value of NSE decreased from 0.961 to 0.891 for the Bao_{hr} model compared to the Bao_{dd} model, and both models performed mostly similarly. Here, for the Bao_{hr} model, we also focus on a site-specific example at a site (DE-Hai) in central Germany with a deciduous broadleaf forest where the Bao_{hr} model proved to be capable of simulating annual average of GPP_{EC} flux relatively well when the model was parameterized for each site–year and each site (Figure 8). However, GPP_{EC} was underestimated in cases of PFT-specific and global parameterization. For this specific site, the Bao_{hr} model performed relatively better in comparison to Bao_{dd} model (Figures 8 and S7 of Supporting Information S1).

3.5. Role of Parameterization Strategies on Model Performance

Model performances at an annual scale generally increased with a more detailed parameterization strategy (Figure 9). For the P_{hr}^{W} model, the median differences in annual model performance between the most detailed







Figure 5. Scatter plot of the annual average (from good quality site–years, see Appendix C) of eddy covariance measurements derived gross primary production (GPP_{EC}) versus simulated gross primary production (GPP_{sim}) from P-model of Mengoli et al. (2022) parameterized at hourly scale (a) without drought stress (P_{hr} model) and (b) with drought stress (P_{hr}^W model). The results in this plot are from parameterization for each site–year. We only used good–quality site–years in this figure (Appendix C). The dots in the scatter represent a site–year and are colored by the aridity index (AI) of the site. The model performance metrics (Nash-Sutcliffe efficiency, viz. NSE) are shown at the top of each subplot. The equations of fitted regression lines are shown in respective subplots.

parameterization strategy, that is, site–year parameterization, and other detailed parameterization strategies, that is, per site parameterization using $Cost^{IAV}$ and per site parameterization were small, which were 0.12, and 0.11, respectively. In contrast, the median differences in annual model performance between the most detailed parameterization strategy, that is, site–year parameterization, and other generalized parameterization strategies, that is, PFT-specific parameterization, and global parameterization were quite large, which were 0.28, and 0.37, respectively. Similarly, for the Bao_{hr} model, the median differences in annual model performance between the most detailed parameterization strategy, that is, site–year parameterization, and other detailed parameterization strategies, that is, per site parameterization using $Cost^{IAV}$ and per site parameterization were 0.16, and 0.15, respectively. In contrast, the median differences in annual model performance between the most detailed parameterization strategy, that is, site–year parameterization, and other detailed parameterization strategies, that is, per site parameterization using $Cost^{IAV}$ and per site parameterization were 0.16, and 0.15, respectively. In contrast, the median differences in annual model performance between the most detailed parameterization strategy, that is, site–year parameterization, and other generalized parameterization strategies, that is, PFT-specific parameterization, and global parameterization were 0.34, and 0.39, respectively. The positive values of median annual model performance confirm the highest median performance of site–year parameterization compared to the other four parameterization strategies.

At an hourly scale, differences in model performance between a pair of similar parameterization strategies, such as a pair of detailed parameterization (i.e., between site–year-specific and site-specific) or a pair of generalized parameterization (i.e., between per PFT and global) approaches for both models were small (Figure S8 of Supporting Information S1). However, this difference can be higher between a detailed and a generalized model parameterization strategy. The median differences in hourly NNSE between site–year-specific and site-specific model parameterization were 0.02 and 0.01 for the P_{hr}^W model and Bao_{hr} model, respectively. In contrast, the median differences in hourly NNSE between site–year-specific and global model parameterization were 0.11 for both P_{hr}^W and Bao_{hr} models.

The median differences in annual model performance between per–site parameterization using $Cost^{IAV}$ and per– site parameterization were relatively small, which were 0.01 and 0.00 for P_{hr}^W model and Bao_{hr} model, respectively, and it shows the additional constraint on IAV of GPP flux in the cost function did not substantially improve annual model performance. At hourly scale, median differences in model performance between per–site





Figure 6. Comparison of annual average of gross primary production (GPP) derived by eddy covariance measurements (GPP_{EC}) , and GPP simulated (GPP_{sim}) by the P-model of Mengoli et al. (2022) parameterized at hourly scale without drought stress (P_{hr} model) and with drought stress (P_{hr}^{W} model). The five subplots show simulated GPP from (a) site–year specific parameterization, (b) site-specific parameterization using an additional constraint on inter–annual variability in the cost function ($Cost^{IAV}$), (c) site-specific parameterization, (d) plant–functional types (PFT) specific parameterization, and (e) global parameterization. The values of model performance measures (Nash-Sutcliffe efficiency, viz. NSE, correlation coefficient, viz. r, relative variability, viz. α_{NSE} , and bias, viz. β_n) are shown on top of respective subplots. This site is dominated by Mulga (*Acacia aneura*), and had an annual average temperature of $\approx 22 \,^{\circ}$ C, and an annual average precipitation of $\approx 318 \,$ mm during the observation period (Cleverly et al., 2013; Pastorello et al., 2020). The vertical error bars represent the uncertainty range (see Section 2.4.3) of the annual average of GPP_{EC} for the corresponding years (the uncertainty range is relatively small for this specific site). The site ID, PFT, and Köppen–Geiger climate class (KG) of the site are provided on top of the figure in bold.

parameterization using $Cost^{IAV}$ and per–site parameterization were also non-existent, which were 0.00 for both P_{hr}^{W} and Bao_{hr} models. Though the per–site parameterization using $Cost^{IAV}$ did not improve the annual model performance, it also did not degrade the hourly model performance.

3.6. Differences Between Modeling Experiments of Various Complexities

The lowest AIC values were obtained for per site–year parameterization for all the models at hourly and daily scales or aggregations, suggesting the sum of squares errors (*SSE*) was substantially reduced even when a comparatively complex parameterization strategy with a large number of model parameters was chosen (Table 3). The AIC values gradually increased from per site–year, per site, per PFT to global parameterization at hourly and daily scales or aggregations for all three models which are P_{hr}^W model, Bao_{hr} model, and Bao_{dd} model (Table 3). Semi-empirical models, that is, Bao_{hr} model and Bao_{dd} model also had mostly lower values of AIC compared to optimality-based P_{hr}^W model even though more parameters were parameterization experiments due to the parameterization at daily scale. Whereas, for the other two models, parameterization and forward runs were performed at an hourly scale and then simulations were aggregated to daily resolution. The P_{hr} model was not included in AIC or AIC_c analysis as previous results proved this model significantly underperformed compared to the other models, and this will always result in higher AIC or AIC_c values.

At monthly and annual scales, we show the differences in AIC_c values between P_{hr}^W , Bao_{hr} , and Bao_{dd} models for the same parameterization strategy, and not between parameterization strategies in a same model. The reason behind this is AIC_c values largely depend on the relationship between sample size, that is, *n*, and the total number of parameters which were parameterized, that is, *K*. The values of AIC_c became very large even when a



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Figure 7. Scatter plot of annual average (from good quality site–years, see Appendix C) eddy covariance derived gross primary production (GPP_{EC}) versus simulated gross primary production (GPP_{sim}) by the light use efficiency model of Bao, Wutzler, et al. (2022) parameterized at hourly (Bao_{hr} model) and daily scale (Bao_{dd} model) for each site–year. The plots show the performance of the (a) Bao_{hr} model, and (b) Bao_{dd} model. The dots in the scatter represent a site–year. The model performance metrics (Nash-Sutcliffe efficiency, viz. NSE) are shown on the top of each subplot. The equations of fitted regression lines are shown in respective subplots.

significantly smaller *SSE* was obtained, and they became unreliable when the value of *n* was closer to *K*. For example, at monthly aggregation, per site–year parameterization of the P_{hr}^W model had a very high AIC_c value of 1.30×10^6 even when it had the lowest *SSE* among all the five parameterization strategies (Tables S3, S4, and S5 of Supporting Information S1). The Bao_{dd} model proved to be better able to capture the seasonal cycle, that is, monthly GPP estimates compared to the other two models for most of the parameterization experiments considering the number of parameterization using *Cost^{IAV}* and per site parameterization at monthly aggregation. In contrast, at an annual scale, the P_{hr}^W model had mostly the lowest AIC_c values, and some of the experiments also suffered from the above-described unreliable AIC_c estimates, where *n* and *K* had similar values (Tables 3, S3, S4, and S5 of Supporting Information S1).

3.7. Model Performances Across Different Plant-Functional Types

Generally better model performances were achieved with both the P_{hr}^W model and the Bao_{hr} model when parameterized detailed model parameterization strategies were used (Figure 10). In this analysis, we removed the per–site parameterization experiment using $Cost^{IAV}$ as it performed very similar to per–site parameterization, and also we did not consider the P_{hr} model as it produced poor performance across all the PFTs.

The highest median NNSEs were obtained with per–site–year parameterization for almost all the PFTs for both models. The only exception is WSA for which the median model performance (median NNSE: 0.85) was marginally higher for per–site parameterization experiment using $Cost^{IAV}$ compared to per–site–year parameterization (median NNSE: 0.84) of Bao_{hr} model. For the P_{hr}^W model parameterization experiments, the highest median value of NNSE was found for CSH for per–site–year parameterization (median NNSE: 0.88), DBF for per–site parameterization (median NNSE: 0.85), CSH and DBF for per–PFT parameterization (median NNSE: 0.81), and CSH for global parameterization (median NNSE: 0.80). However, CSH had only three sites and highest median model performance for CSH should be interpreted with caution. For the Bao_{hr} model parameterization (median NNSE: 0.89), DBF for per–site–year parameterization (median NNSE: 0.89), DBF for per–Site parameterization (median NNSE: 0.87), DBF for per–Site par





Hainich (Site ID: DE-Hai PFT: DBF KG: Cfb)

Figure 8. Comparison of annual average of gross primary production (GPP) derived by eddy covariance measurements (GPP_{EC}) , and GPP simulated (GPP_{sim}) by the light use efficiency model of Bao, Wutzler, et al. (2022), which was parameterized with hourly data (Bao_{hr} model). The five subplots show simulated GPP from (a) site–year specific parameterization, (b) site-specific parameterization using an additional constraint on inter–annual variability in the cost function $(Cost^{IAV})$, (c) site-specific parameterization, (d) plant–functional types (PFT) specific parameterization, and (e) global parameterization. The years 2010–2012 could not be parameterized in the case of site–year parameterization, as there were no good quality evapotranspiration estimates from latent heat flux measurements for those years. The values of model performance measures (Nash-Sutcliffe efficiency, viz. NSE, correlation coefficient, viz. *r*, relative variability, viz. α_{NSE} , and bias, viz. β_n) are shown on top of respective subplots. This site represents an average 140-year-old deciduous forest (Tamrakar et al., 2018) with a distinct seasonal cycle and an annual average temperature of ≈ 8.3 °C, and an annual average precipitation of 750–800 mm during the observation period (Knohl et al., 2003a; Pastorello et al., 2020). The vertical error bars represent the uncertainty range (see Section 2.4.3) of the annual average of GPP_{EC} for the corresponding years. The site ID, PFT, and Köppen-Geiger climate class (KG) of the site are provided on top of the figure in bold.

parameterization (median NNSE: 0.82), and DBF and MF for global parameterization (median NNSE: 0.81). We found similar results also for climate–vegetation types (Section 2.2), where a more detailed parameterization strategy achieved higher model performance than a generalized parameterization strategy (Text S5 and Figure S9 of Supporting Information S1).

Specifically, we found the model performance significantly differed between per–site–year parameterization and global model parameterization for most PFTs and both models (Figure 10). Model performance was also similar between a pair of detailed parameterization strategies (per–site–year and per–site parameterization) and general parameterization strategies (per-PFT and global parameterization) for most PFTs. For PFTs with very few sites, such as CSH, DNF, SNO, and WSA, statistical significance testing showed that the model performance was similar among parameterization strategies. However, statistical significance tests can be unreliable for a few samples. The model performance between P_{hr}^W model and the Bao_{hr} model for a specific parameterization strategy was similar for most PFTs (Table S6 of Supporting Information S1). Similarly, model performance between both models for a specific parameterization strategy was also identical for most PFTs (Table S7 of Supporting Information S1).

3.8. Correlation Between Annual Model Performance and Model Performance in Simulating Diurnal GPP Peaks

One of the crucial reasons behind poor annual model performance (Figure 2 and Table D1) can be the inability of both the P_{hr}^{W} model and the Bao_{hr} model to capture the peaks of GPP_{EC} (Figure S10 of Supporting

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Figure 9. Distributions of the differences between model performance measures (normalized Nash-Sutcliffe efficiency, viz. NNSE) calculated at annual scale, from various pairs of model parameterization experiments conducted for the P-model of Mengoli et al. (2022) with drought stress, parameterized with hourly data (P_{hr}^W model) and the light use efficiency model of Bao, Wutzler, et al. (2022), parameterized with hourly data (Bao_{hr} model). *Cost^{IAV}* in parentheses denotes the usage of an additional constraint on annual gross primary production flux during per–site parameterization. The boxes are spanned between the first and third quartiles of the differences, and the line in the middle represents the median. The whiskers show the farthest data point from the box within 1.5× of the interquartile range. The circles represent the outliers that go beyond the limits of the whiskers. The vertical dotted gray lines separate each pair of model parameterization strategies.

Information S1). Specifically, $P90_{GPP_{EC}}$ was highly underestimated in the case of global parameterization. The median of the ratio of $P90_{GPP_{sim}}$ to $P90_{GPP_{EC}}$ were 0.77 and 0.84 for the P_{hr}^W model and the Bao_{hr} model, respectively, during global parameterization. The underestimation generally decreased with more detailed parameterization strategies, with little difference between the models. The median values of the ratio of $P90_{GPP_{sim}}$ to $P90_{GPP_{EC}}$ were 0.95 and 0.93 for the site–year parameterization of the P_{hr}^W model and the Bao_{hr} model, respectively. Moreover, the lower values of the interquartile range (IQR) of these ratios signify the importance of site–year parameterization compared to per PFT or global parameterization to reliably capture the peak GPP_{EC} in diurnal cycles for most of the sites and to attain better model in the case of site–year parameterization, 0.44 and 0.37 for the P_{hr}^W model and the Bao_{hr} model, respectively in the case of PFT-specific parameterization, and 0.44 and 0.49 for the P_{hr}^W model and the Bao_{hr} model, respectively in the case of global parameterization.

We further found that if a certain parameterization strategy better simulated the $P90_{GPP_{EC}}$ for each site–year, it corresponded to a comparatively better annual model performance for a site which is demonstrated by the positive values of Pearson correlation coefficients (Figure 11). Here also, a detailed parameterization strategy, such as site–year parameterization resulted in a better simulation of $P90_{GPP_{EC}}$, and thus better annual model performance for most of the sites compared to a generalized parameterization strategy, such as global parameterization. In this case when *j*1 was site–year parameterization and *j*2 was global parameterization, 91% and 89% sites had higher $NNSE_{P90}^{j1}$ than $NNSE_{P90}^{j2}$ and corresponding $NNSE_{y}^{j1}$ than $NNSE_{y}^{j2}$ for the P_{hr}^{W} model and the Bao_{hr} model, respectively. When *j*1 was parameterization per site using $Cost^{IAV}$ and *j*2 was parameterization per site, respectively, only 34% and 37% had positive values of $\Delta NNSE_{P90}$ (i.e., $NNSE_{P90}^{j0} > NNSE_{P90}^{j2}$) and corresponding



Table 3

Akaike's Information Criterion (AIC) or Corrected AIC (AIC,) Values for Modeling Experiments of Various Complexities

		Parameterization strategies				
Temporal scale/ aggregation	Models	Per site-year	Per site using Cost ^{IAV}	Per site	Per PFT	Global
Hourly scale (AIC)	$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	1.72×10^{7}	1.84×10^{7}	1.86×10^{7}	2.16×10^{7}	2.25×10^{7}
	Bao _{hr}	1.58×10^{7}	1.70×10^{7}	1.68×10^{7}	2.10×10^7	2.16×10^{7}
Daily scale/ aggregation (AIC)	$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	4.58×10^{5}	5.05×10^5	5.11×10^{5}	6.88×10^5	7.42×10^{5}
	Bao _{hr}	3.99×10^{5}	4.28×10^5	4.24×10^{5}	6.65×10^5	6.92×10^{5}
	Bao _{dd}	2.68×10^{5}	3.12×10^5	3.18×10^{5}	4.98×10^5	5.58×10^{5}
Monthly aggregation (AIC_c)	$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	1.30×10^{6}	1.63×10^4	1.63×10^4	2.04×10^4	2.25×10^4
	Bao _{hr}	-4.55×10^4	1.79×10^4	1.79×10^{4}	1.97×10^4	2.02×10^4
	Bao _{dd}	-3.40×10^4	1.67×10^4	1.70×10^4	1.52×10^4	1.73×10^4
Annual aggregation (AIC _c)	$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	-3.56×10^{3}	-5.22×10^3	-5.57×10^{3}	9.09×10^2	9.28×10^2
	Bao _{hr}	-3.19×10^{3}	-3.84×10^{3}	-3.72×10^{3}	1.42×10^{3}	9.26×10^2
	Bao _{dd}	-3.66×10^{3}	-3.45×10^3	-3.39×10^{3}	1.12×10^{3}	7.49×10^2

 $\Delta NNSE_y$ (i.e., $NNSE_y^{j1} > NNSE_y^{j2}$) for the P_{hr}^W model and the Bao_{hr} model, respectively. This signified that using an additional constraint related to the IAV of GPP in the cost function during model parameterization did not improve the prediction of peak GPP values for most of the sites.

4. Discussion

4.1. Uncertainties in Modeling Experiments

Any model-data-integration study is prone to uncertainties related to both data and the model. The EC data set used in our study has a couple of well-known uncertainties. For example, the NEE measurements using the eddy covariance technique can have uncertainties due to the accumulation of atmospheric CO₂ under the canopy at night (storage) and a sudden turbulent mixing during the morning when the stable night-time boundary layer breaks up, or because of advection of atmospheric CO₂ out of the control volume sampled by the eddy covariance system (Aubinet, 2008; D. Baldocchi et al., 2000; Jocher et al., 2018). The GPP fluxes that we used were derived from NEE measurements by extrapolating the night-time respiration of the ecosystem (Reichstein et al., 2005) to daytime. Moreover, GPP can be estimated based on another well-known algorithm, the daytime partitioning method (Lasslop et al., 2010). We preferred night-time partitioning as only respiration is modeled in this method. In daytime partitioning, both GPP and respiration are modeled, resulting in higher prediction errors. The uncertainties in our modeling results due to the choice of partitioning algorithm should be small as quantified in a previous study by Desai, et al. (2008). Papale et al. (2006) proposed a standardized set of flux correction techniques for the above-described issues related to EC methods and removing random spikes in EC measurements. Many of these correction techniques were adapted to the standardized ONEFlux pipeline for the production of the FLUXNET2015 data set (Pastorello et al., 2020). Papale et al. (2006) also found that the uncertainties associated with annual NEE, and fluxes derived from NEE, such as GPP and TER have an inherent uncertainty well below 100 gC \cdot m⁻² \cdot year⁻¹. In our study, we found that annual *GPP*_{Sim} produced by experiments related to detailed parameterization strategies were mostly within the range of uncertainty of GPP_{EC} on the choice of different variables produced by the data processing pipeline. However, quantifying full uncertainty in GPP_{EC} related to measurement errors, random noises, data gaps and gap-filling methods, flux correction methods, and flux partitioning methods as well as propagating all the uncertainties to various temporal scales is a very challenging task and beyond the scope of this study. Similarly, GPP_{Sim} can also have various uncertainties related to model structure and simplified representation of ecosystem functions (such as stomatal conductance, photosynthesis, leaf energy balance etc.), forcing data, and parameters (Schaefer et al., 2012; Zheng et al., 2018). For example, the Bao model used in our study is based on the big leaf assumption. In contrast, differentiating between sunlit and shaded leaves (two big-leaf assumption) in a similar model can lead to higher performance under certain conditions, such as under hot and dry conditions at daily and weekly scales (Bao, Ibrom, et al., 2022). However, the two big-leaf model was also equally poor at explaining IAV. In this study, we aimed to address some of these



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Figure 10. Box-plots showing the range of the hourly model performance metric (normalized Nash-Sutcliffe efficiency, viz. NNSE), for the sites in different plant–functional types (PFT), and different parameterization experiments. The subplots show the model performance for (a) P-model of Mengoli et al. (2022) with drought stress function, parameterized with hourly data (P_{hr}^{W} model), and (b) the light use efficiency model of Bao, Wutzler, et al. (2022) parameterized with hourly data (Bao_{hr} model). The numbers in parentheses beside the name of each PFT on the *x*-axis are the number of sites present in a specific PFT. The boxes are spanned between the first and third quartiles of NNSE values, and the line in the middle represents the median. The whiskers show the farthest data point from the box within 1.5× of the interquartile range. The circles represent the outliers that go beyond the limits of the whiskers. The results of statistical significance testing using a two-sample Kolmogorov-Smirnov test (Hodges, 1958) between a pair of model parameterization strategies (connecting bars over boxplots) are shown as * (0.05 > *p*-value ≥ 0.01), ** (p-value ≥ 0.001), *** (p-value < 0.001) if the distributions of model performances were not identical. No bars and star symbols over a pair of boxplots signify that the distributions of model performances from a pair of parameterization strategies were identical and the null hypothesis could not be rejected (p-value ≥ 0.05). For, deciduous needle-leaf forests (DNF), and areas covered by snow (SNO) only the median value could be shown as these PFTs have only one site. The vertical dotted gray lines separate each PFT.

uncertainties in GPP_{Sim} by adopting majorly two different model structures and various parameterization schemes.

We also used ET in the cost function which is equivalent to latent heat flux. The mismatch between the summation of latent, sensible, and ground heat fluxes with net radiation calculated using incoming and outgoing radiation, the so-called lack of energy–balance closure, remains a long-standing challenge with EC measurements (Foken, 2008; Mauder et al., 2020; Zhang et al., 2024). Quality control of millions of data points at an hourly scale was also challenging, especially when we merged data from various sources, such as in-situ measurements, modeled re-analysis data, and remote sensing-based estimates. Another major uncertainty arises from the mismatch between the footprint of EC towers and the grid of remote sensing data which were used to calculate vegetation indices (Chu et al., 2021). The PFT classification of sites based on a simple PFT classification method may not accurately represent the vegetation of some of the sites. For example, a site at Alice Springs (AU-ASM) in central Australia was classified as a savanna in FLUXNET2015 (Pastorello et al., 2020). In fact, this site is dominated by a discontinuous canopy of Mulga (*Acacia aneura*) that has needle leaves and a seasonal understory grassy layer (Cleverly et al., 2013). This site can be classified as a woody savanna as well. An arctic site in Bayelva (SJ-Blv) has a combination of snow, wet grounds, and specific tundra vegetation (Boike et al., 2018) which were not well represented by the snow classification of FLUXNET2015 (Pastorello et al., 2020). Another limitation of the data set is that sites are mostly clustered in European and North American countries and, hence do





Figure 11. Scatter between differences in model performance in simulating peak gross primary production, viz. GPP $(\Delta NNSE_{P90})$ and annual average of GPP $(\Delta NNSE_y)$. *j*1 and *j*2 are a pair of parameterization strategies for which differences are calculated in each subplot from (a to j). Each dot in the scatter represents a site. The P_{hr}^{W} model and Bao_{hr} model are P-model of Mengoli et al. (2022) with drought stress function and the light use efficiency model of Bao, Wutzler, et al. (2022), both were parameterized with hourly data. The values on top of each subplot indicate the Pearson correlation coefficient (Kirch, 2008) for respective models.

not necessarily represent global ecosystem functioning particularly well due to sampling bias (Papale et al., 2015). Similarly, some PFTs are represented by very few sites, which makes PFT-specific parameterization challenging.

4.2. General Performance of Models in Simulating GPP

Mengoli et al. (2022) evaluated the P_{hr} model across 10 sites consisting of boreal forest, temperate deciduous broadleaf forest, mixed forest, tropical forest, and temperate grassland. They found reasonable performance of P_{hr} model for these sites. In this study, we extended the evaluation of optimality-based P_{hr}^W and P_{hr} models across a wide range of sites, representing various vegetation and climate types. We uncovered poor model performance of the P_{hr} model at many sites, especially at arid sites. Calculating *WAI* using a simple hydrological model and

inducing a moisture stress function, that is, the introduction of the P_{hr}^W model substantially improved model performance to simulate the annual average of GPP fluxes across many sites, including both water–limited and energy–limited sites. Inclusion of the moisture stress function in the P_{hr} model improved not only the annual model performance but also the model performance across all the temporal scales or aggregation levels. This highlights the importance of representing soil moisture conditions in modeling approaches, which aim to accurately represent ecosystem functioning and vegetation response. Our findings also support another study by Schaefer et al. (2012) which found mostly poor performance of 26 models in simulating GPP under dry conditions across EC sites in North America. Even though the drought stress function used in this study does not explicitly account for long legacy effects, our results still confirmed the substantial effect of drought stress on GPP as found by numerous other studies (Anderegg et al., 2015; Müller & Bahn, 2022; Z. Yu et al., 2017; X. Yu et al., 2022). However, the coupling of the hydrological model raised the need to calibrate nine more parameters, which counters the vision of developing a parameter–sparse approach using theories that demand a lower site or site–year specific fine–tuning of model parameters (Prentice et al., 2015). Further experimentation is needed to find a balance between the number of key model parameters, which require calibration, and an accurate representation of ecosystem processes.

Coming to the differences in model structure, we found that semi-empirical models (Bao_{hr} and Bao_{dd} models) performed statistically better, that is, had a lower value of AIC compared to optimality-based model (P_{hr}^{W} model) at hourly and daily scale or aggregations for most of the parameterization experiments even though the semi-empirical modeling experiments needed more parameters to be parameterized. At the monthly aggregation level, the seasonal cycles were also significantly better captured by the parameter-heavy semi-empirical model parameterized with daily data (Bao_{dd} model) for most of the parameterization experiments. However, at the annual aggregation level, the optimality-based model, that is, the P_{hr}^{W} model was comparatively better in most cases and a more flexible semi-empirical model with a higher number of parameters did not have a substantial improvement in annual model performance. In an earlier study, Sims et al. (2005) found that midday gross CO₂ flux is highly correlated with daily and 8-day average of gross CO₂ flux and the midday values can be used to estimate fluxes at higher timescales, that is, at daily and 8-day. In this study, we also found that both hourly and daily GPP estimates from the Bao_{hr} model, Bao_{dd} model, respectively can be used to estimate fluxes at weekly and monthly scales reliably. However, both models poorly estimated the annual average GPP flux per site.

Though the partial sensitivity functions of environmental variables used in the Bao_{hr} model and the Bao_{dd} model were found to be applicable for most of the sites, they can be of many different types and may vary across site conditions (Bao, Wutzler, et al., 2022). The EC sites were also affected by human management, such as irrigation, harvesting, and mowing as well as natural disturbances, such as fire, and pest attacks. These factors can affect the IAV of GPP flux which was estimated from EC measurements. Models used in this study may not be able to account for all of these factors due to structural limitations. For example, in the hydrological model, we only used precipitation and ET to calculate the mass balance of water. However, human management (such as irrigation and drainage) can play an important role, and the *WAI* estimates in managed sites, such as at an irrigated maize site (US-Ne1) may not be accurate.

4.3. The Importance of the Parameterization Approach on Estimating IAV of GPP

We also performed inter–comparison between five different parameterization approaches for both an optimalitybased model and a semi-empirical model, and also for a large number of sites. We did these extensive evaluations to emphasize the importance of model parameterization in capturing the IAV of GPP and gain higher confidence in our conclusions. Model parameterization largely determined model performances and calibrated parameters captured the individual characteristics of sites or climatic events of the site–years (J. Wu et al., 2012). Earlier studies (Groenendijk et al., 2011; Huang et al., 2021) found higher variability of model parameters across sites within a PFT and detailed parameterization was necessary to better simulate carbon fluxes as there can be variations in vegetation functioning and climatic conditions between sites and a simple PFT classification may not be enough to account for these variations. Similarly, we also found detailed model parameterization strategies, such as parameterization specific to site–years or sites comparatively better predicted the annual average of GPP fluxes and year-specific parameters explained some parts of the IAV of GPP flux. At hourly and daily scales, we found slightly better model performance in the case of site–year parameterization compared to site-specific parameterization. This can mainly be due to variations in hydrological conditions, such as drought or excess rainfall between years in a site and site-year-specific parameters may be able to better capture these events. However, in the annual scale model performance was similar between site-year and site-specific parameterization, even though we assumed parametrizing models using data from each site-year would be advantageous to capture IAV of GPP. Changes in parameters between years could reflect structural model issues, such as changes in state variables not represented correctly in the model (e.g., soil moisture) or absent in the model, but with a possible role on IAV (e.g., reserve or enzymatic pools, other biogeochemical cycles, such as nitrogen, or biotic factors promoting, e.g. herbivory). Changes in parameters could also reflect observation artifacts, such as biases, or changes in instrumentation. We would argue that the difference from site to site-year would reflect possible enhancements in performance via improvements in model structure, or via statistically learning temporal dynamics of parameters, which would add an inference component from the influence of environmental forcing on the responses from missing processes. Moreover, as the fast rate of change in climatic characteristics has become more frequent in recent years, developing a generalized model structure to simulate carbon fluxes between years and/or between sites of similar vegetation types has become even more challenging (Knauer et al., 2023).

The generalized model parameterization strategy, that is, global parameterization was also dominated by PFTs, such as ENF and GRA which were represented by many sites, and certain PFTs, such as DNF was represented by only one site in the FLUXNET2015 data set (Pastorello et al., 2020). This may imply that global parameterization or parameter up-scaling experiments using the FLUXNET2015 data set (Pastorello et al., 2020) may result in biased parameter sets that cannot be generalized to the global scale or a weighted site representation may be necessary in this case. Besides model parameterization, a recent study by Zou et al. (2024) highlighted that the importance of each independent driver, and their relative contributions vary over time and ecosystem type. The relative importance of forcing variables may also be another factor besides model parameterization. We also found both the P_{hr}^W model and the Bao_{hr} model showed the highest model performance at the sub-daily scale mostly for forest sites compared to savannas or grasslands, this in turn led to the poor simulation of IAV at many sites which are not forests. In most cases, we found that parameterization strategy played an important role in the variability of model performance across PFTs. For most PFTs, model performance between per–site–year and global parameterization were statistically different. However, there was no significant difference between model performance across PFTs between P_{hr}^W model and Bao_{hr} model.

Though we have demonstrated the capability of the P_{hr}^{W} model that included drought stress and the Bao_{hr} model to simulate the hourly fluxes of GPP, accurate estimation of IAV of GPP fluxes at the site level with these models requires further developments. Particularly, both models failed to capture the peak GPP in diurnal cycles at many sites even after model parameterization at a sub-daily scale and using an additional constraint on the IAV of GPP in the cost function. These underestimations at an hourly scale may have accumulated to a larger error when the fluxes were aggregated at an annual scale to study the IAV. We also showed that comparatively better model performances were achieved when the GPP peaks per site-year were better simulated. These results are similar to another study by Lin et al. (2023), directed at evaluating terrestrial ecosystem models' capability in explaining the IAV of GPP which also found an underestimation of GPP. Though the occurrences of peak GPP in a diurnal cycle can be comparatively smaller, an earlier study by Zscheischler et al. (2016) found that high GPP values during ideal weather conditions majorly contributed to the IAV of GPP flux. While the study of Zscheischler et al. (2016) was performed for temperate forests, we found a similar pattern for sites representing diverse climate-vegetation types in our study. Paschalis et al. (2015) also found that short-term variability of meteorological forcing can affect carbon and water fluxes from hourly to annual or even at longer scales, especially if they affect soil water availability and similar slow processes. Another study by Fatichi and Ivanov (2014) which involved synthetic data analysis also concluded that random occurrences of favorable weather conditions and subsequent higher carbon assimilation explain IAV of aboveground NPP better than average weather conditions of a year or growing season. Unlike some of the previous studies mentioned above which were mainly focused on finding correlations between GPP peaks and annual GPP from either only observed data or synthetic data, here, we found a direct correlation between model performance in simulating GPP peaks and model performance in simulating IAV of GPP. It is also true that some of the peak values found in EC-derived diurnal GPP can also be outliers, which are artifacts of data processing algorithms such as the gap-filling algorithm. It can be hard to distinguish these outliers from the true peaks of GPP. However, the bias introduced by the gap-filling algorithm is relatively low (Moffat et al., 2007). Therefore, it is unlikely that the reasoning behind the general underestimation of GPP peaks by models is due to the bias introduced during gap-filling.



The poor representation of IAV of GPP can be attributed to either limitations of models or model parameterization strategies. It is important to discover which seasonal phases of the GPP dynamics for particular vegetation types or climatic zones are not well represented in models simulating the IAV of GPP. In our study, we accounted for biophysical forcings, such as meteorological variables in our models, and used fAPAR to represent phenological dynamics. However, if there is a decoupling between the phenological and GPP dynamics, the length since GPP onset and senescence can be another factor behind IAV, especially during climate extremes, or large meteorological time shifts. Previous studies such as Wolf, et al. (2016) showed that the net effect of climate extremes, that is, the 2012 summer drought on carbon uptake was compensated by warmer springs, and increased soil moisture consumption during warmer springs led to the drought conditions in summer. Similar findings were also reported by van der Woude et al. (2023) for European forests where the effect of the 2018 drought was offset by prolonged carbon uptake during warm autumn. Bastos et al. (2020) and Smith et al. (2020) reported a similar finding that the increased vegetation growth during warm spring was responsible for soil moisture depletion and inflation of the 2018 drought effect in Europe. It is particularly important to focus on the meteorological sensitivity of GPP during periods of high productivity where improvements in the prediction of high fluxes would tend to improve the description of IAV. Another aspect could also be to decompose the metric (Gupta et al., 2009) used in the cost function or develop a more detailed model evaluation to understand which other parts of the time series were not well constrained during model parameterization.

5. Conclusions

We have demonstrated the capability of an improved version of an optimality-based model (P_{hr}^{W} model) and a semi-empirical LUE model (Bao_{hr} and Bao_{dd} models) to simulate sub-daily or daily GPP fluxes across 198 EC sites, representing 13 different vegetation types including forests, grasslands, savannas, croplands, and tundra. We also performed various model parameterization strategies to systematically evaluate the factors affecting the IAV of GPP. The main conclusions from our study are:

- We found that the semi-empirical model mostly produced better results at hourly, daily, and monthly scales compared to the optimality-based model. At an annual scale, the improvement in the performance of the semiempirical model was not significant even though more parameters were parameterized to flexibly capture the ecosystem dynamics. Both optimality-based and semi-empirical models performed better on hourly, daily, weekly, and monthly scales compared to annual scale.
- 2. Model structure was an important factor behind variability in both hourly and annual model performances when drought stress was not explicitly accounted for in the optimality-based model. However, differences in climate classes and parameterization strategies became important factors behind variability in hourly and annual model performances, respectively, when the previous version of the optimality-based model (without drought stress) was not considered in the analysis.
- 3. Explicit accounting of drought stress in the optimality-based ecosystem model is a necessity as it proved to be an important factor in controlling GPP fluxes at all temporal scales including at annual aggregation. The explicit representation of drought stress not only improved the performance of the optimality-based model at arid sites but also at energy–limited sites.
- 4. Both models performed better mostly at forest sites compared to grasslands or savannas which may also lead to poor estimation of IAV of GPP at many sites globally.
- 5. While these models generally performed well in simulating hourly GPP dynamics, the small errors at the subdaily scale, particularly related to the estimation of GPP peaks, accumulated to bigger errors at the annual scale and led to poor performance of models in explaining the IAV of GPP. We found that comparatively better annual model performance could be achieved when the peaks of GPP were better simulated.

Our results further suggest the need to focus on sub-daily GPP dynamics during the various seasonal phases, especially highly productive ones, toward an improved constraint on GPP sensitivities. Hence, better annual model performance with a detailed parameterization strategy, such as site–year parameterization, signifies that temporally varying model parameters are necessary to better capture the variations of annual average GPP and indicate that ecosystem functioning is not stable between years. These new understandings can guide us toward developing models and parameterization strategies for simulating the inter–annual variations in ecosystem GPP more successfully, and improve our understanding of the global carbon cycle response to changing climatic conditions.



Appendix A: Data Description

Table A1Description of Force	ing and Model Parameterization Data			
Abbreviation	Definition	Unit	Variable name in data set/remarks	Reference
GPP _{EC} ^a	GPP derived from EC based net ecosystem exchange (NEE) using night-time partitioning [*] method. Used for model calibration and evaluation.	$\mu molCO_2 \cdot m^{-2} \cdot s^{-1}$	GPP_NT_VUT_USTAR50	Pastorello et al. (2020); Reichstein et al. (2005)
GPP_NT_VUT_05	GPP derived from percentile 5 of 40 different estimates of NEE using night-time partitioning method. Used to quantify uncertainty in GPP_{EC} .	$\mu molCO_2 \cdot m^{-2} \cdot s^{-1}$	GPP_NT_VUT_05	Pastorello et al. (2020); Reichstein et al. (2005)
GPP_NT_VUT_95	GPP derived from percentile 95 of 40 different estimates of NEE using night-time partitioning method. Used to quantify uncertainty in <i>GPP_{EC}</i> .	$\mu molCO_2 \cdot m^{-2} \cdot s^{-1}$	GPP_NT_VUT_95	Pastorello et al. (2020); Reichstein et al. (2005)
GPP_DT_VUT_05	GPP derived from percentile 5 of 40 different estimates of NEE using daytime partitioning method. Used to quantify uncertainty in GPP_{EC} .	$\mu molCO_2 \cdot m^{-2} \cdot s^{-1}$	GPP_DT_VUT_05	Pastorello et al. (2020); Lasslop et al. (2010)
GPP_DT_VUT_95	GPP derived from percentile 95 of 40 different estimates of NEE using daytime partitioning method. Used to quantify uncertainty in GPP_{EC} .	$\mu molCO_2 \cdot m^{-2} \cdot s^{-1}$	GPP_DT_VUT_95	Pastorello et al. (2020); Lasslop et al. (2010)
σ_{NEE}	Random uncertainty for NEE	$\mu molCO_2 \cdot m^{-2} \cdot s^{-1}$	NEE_VUT_USTAR50_RANDUNC	Pastorello et al. (2020)
LE	Latent heat flux	${ m W}\cdot{ m m}^{-2}$	LE_F_MDS	Pastorello et al. (2020)
σ_{LE}	Random uncertainty for latent heat flux	${ m W}\cdot{ m m}^{-2}$	LE_RANDUNC	Pastorello et al. (2020)
SW_IN ^b	Incoming shortwave radiation	$W \cdot m^{-2}$	SW_IN_F	Pastorello et al. (2020)
NETRAD ^{b, c}	Net radiation	$W \cdot m^{-2}$	NETRAD	Pastorello et al. (2020)
SW_IN_POT	Potential incoming shortwave radiation	$W \cdot m^{-2}$	SW_IN_POT	Pastorello et al. (2020)
PPFD_IN ^a	Incoming photosynthetic photon flux density	$\mu mol\ photons\cdot m^{-2}\cdot s^{-1}$	<i>PPFD_IN</i> gap-filled with 2.04 × <i>SW_IN</i>	Pastorello et al. (2020); see Section 3.4.2 of Stocker et al. (2020) for the gap-filling equation
$T^{\mathbf{b}}$	Air temperature	°C	TA_F_MDS	Pastorello et al. (2020)
VPD ^b	Vapor pressure deficit	Pa	VPD_F_MDS	Pastorello et al. (2020)
$P^{\mathrm{b,d}}$	Precipitation	$mm\cdot h^{-1}$ or $mm\cdot d^{-1}$	Р	Pastorello et al. (2020)
<i>CO</i> ₂	Atmospheric CO ₂ concentration dry air mole fractions from quasi- continuous measurements at Mauna Loa	ppm	CO ₂ _mlo_surface- insitu_1_ccgg_DailyData (interpolated linearly to hourly scale). The measurements from Mauna Loa were used for all sites as the CO ₂ concentration measurements at EC sites are often noisy and discontinuous.	Thoning et al. (2021)
elev	Site elevation	m a.s.l.	Collected from literature	Bao, Wutzler, et al. (2022)
ET_{LE}^{a}	Evapotranspiration derived from <i>LE</i> flux	$mm \cdot h^{-1}$ or $mm \cdot d^{-1}$	Calculated from LE with a dependency on T	Henderson-Sellers (1984)



Table A1

Continued				
Abbreviation	Definition	Unit	Variable name in data set/remarks	Reference
σ_{ET}	Random uncertainty for ET_LE	$mm \cdot h^{-1}$ or $mm \cdot d^{-1}$	Calculated from <i>LE_RANDUNC</i> with a dependency on <i>T</i>	Henderson-Sellers (1984)
PET	Potential evapotranspiration	$mm \cdot h^{-1}$ or $mm \cdot d^{-1}$	Calculated from <i>T</i> , <i>NETRAD</i> and <i>elev</i> using the method of Priestley and Taylor	Priestley and Taylor (1972)
CI	Cloudiness index	-	Calculated as $1 - \left(\frac{SW_{IN}}{SW_{IN}_{POT}}\right)$	Bao, Wutzler, et al. (2022); Fu and Rich (1999); Turner et al. (2006)
WAI	Water availability indicator	mm	Described in Text S1 of Supporting Information S1	Bao, Wutzler, et al. (2022); Tramontana et al. (2016); Trautmann et al. (2018)
W	Soil water supply	$\mathrm{mm} \cdot \mathrm{mm}^{-1}$	Calculated as $\frac{WAI}{AWC}$ (AWC is defined in Table 1)	Bao, Wutzler, et al. (2022)
NDVI	Normalized difference vegetation index	_	Daily <i>NDVI</i> from FluxnetEO v2 (MODIS) was linearly interpolated to hourly	Walther et al. (2022, 2023)
fAPAR	Fraction of incident photosynthetic photon flux that is absorbed by vegetation	-	Linear relationship between <i>NDVI</i> and <i>fAPAR</i> was assumed. $\begin{cases} NDVI, & \text{if } NDVI > 0\\ 0, & \text{if } NDVI \le 0 \end{cases}$	Bao, Wutzler, et al. (2022); Myneni et al. (1997)
QC ^a	Data quality flags	_	1.0 (good quality), 0.5 (medium quality), and 0.0 (bad quality) in the case of hourly data, which is the fraction of good quality measured or gap-filled data from two half-hours. In the case of daily, QC can have any values between 0.0 and 1.0, which is a fraction representing the percentage of good quality measured or gap- filled data in a day. The daily data with $QC > 0.8$ was considered good	Pastorello et al. (2020); Nelson et al. (2024)

Note. *We preferred the night-time partitioning (Reichstein et al., 2005) over daytime partitioning (Lasslop et al., 2010) as only respiration is modeled in this case and GPP is derived as the difference between measured NEE and respiration. Whereas, in the daytime partitioning method, GPP is modeled as well and can have prediction errors due to uncertain model parameters. ^aFor GPP_{EC} , the QC flags of *NEE*, and for ET_{LE} the QC flags of *LE* were used, as they were derived from the respective variables. *QC* flags of *SW_IN* were used to determine bad and medium quality data of *PPFD_IN*, which were replaced with a gap-filling procedure. ^bBad, medium quality (value of *QC* is 0 and 0.5) data and gaps were filled with downscaled (Besnard et al., 2019) ERA5 (Hersbach et al., 2023) or ERA-Interim v2.0 data (Berrisford et al., 2011). ^cWe have collected good quality *SW_IN* and *NETRAD* values from all the sites and fitted a linear regression model using the RANdom SAmple Consensus (RANSAC) algorithm (Fischler & Bolles, 1981) to determine the relation between them. The fitted equation (*NETRAD* = 0.7066 × *SW_IN* - 0.1345) was used to fill gaps in *NETRAD* using *SW_IN*. The gap-filling with regression was only applied for a few sites at hourly scale. ^dAt hourly scale, the data gaps or bad quality data in *P* were filled by distributing the daily downscaled *P* (Besnard et al., 2019) from ERA-Interim v2.0 (Berrisford et al., 2011) for a certain day to the hourly timesteps, based on hourly *P* from gridded ERA5 data (Hersbach et al., 2023).

Appendix B: Data Screening for Model Parameterization

We used only good–quality data to calibrate model parameters. At hourly scale, we selected GPP_{EC} and ET_{LE} as good quality data when the values of their respective QC flag were 1 (Table A1). At the daily scale, we considered a GPP_{EC} and ET_{LE} data point as good when the value of the QC flag was greater than 0.8. We also removed any data gaps from observed and simulated data, σ_{NEE} , and σ_{LE} (Table A1). There were certain negative values in our GPP_{EC} data, as it was calculated using night-time based partitioning method (Reichstein et al., 2005). In this case, if a negative GPP_{EC} value occurred, when the SW_{LIN} (Table A1) is zero that is, during night hours, we replaced



those data points with 0 and used them in the cost function. If the negative GPP_{EC} occurred during day hours, we excluded them.

Appendix C: Data Screening for Evaluation of Model Performance

The good quality data at an hourly scale were selected using the same criteria described in Appendix B. The data screening at a daily scale was also similar to Appendix B, when the LUE model was parameterized using daily data. For all other cases, we assigned a flag (0 = not considered, 1 = considered) to identify which data points were considered during model parameterization. We aggregated this flag to daily, weekly, monthly, and annual scales by taking averages. Then this flag indicated the fraction of good quality data used to calculate a data point in a certain temporal resolution. We only used data points at certain temporal resolutions which were calculated using more than 50% (flag value > 0.5) good quality data points from hourly/daily resolution. We calculated monthly, and annual model performance metrics for a certain site if at least three data points were present. We couldn't calculate annual metrics for 76 and 85 sites due to low numbers of good quality site–years when the annual data was aggregated from hourly, and daily data, respectively. The monthly metrics were not calculated for the three sites due to the same reason when they were aggregated from daily data.

Appendix D: Median Values of Model Performance

The median values of the model performance metric, that is, NNSE which are plotted in Figure 2 are summarized in Table D1.

	-	·				
		Parameterization strategies				
Temporal scale/ aggregation	Models	Per site-year	Per site using Cost ^{IAV}	Per site	Per PFT	Globa
Hourly/ daily scale	$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	0.827	0.799	0.805	0.738	0.712
	$\mathbf{P}_{\mathbf{hr}}$	0.478	0.470	0.469	0.461	0.490
	Bao _{hr}	0.855	0.837	0.838	0.758	0.730
	Bao _{dd}	0.837	0.796	0.796	0.686	0.642
Annual aggregation	$\mathbf{P}_{\mathrm{hr}}^{\mathrm{W}}$	0.543	0.405	0.373	0.201	0.143
	P_{hr}	0.018	0.019	0.018	0.018	0.019
	Bao _{hr}	0.628	0.460	0.460	0.187	0.189
	Bao _{dd}	0.704	0.449	0.475	0.233	0.151

Table D1

Median NNSE Obtained at Each Modeling Experiment at Hourly/Daily Scale and Annual Aggregations

Appendix E: Mean Values of Fraction of Site–Years per Site for Which Simulated GPP Was Within Uncertainty Range

The mean values of fractions of site-years per site for which the annual average of GPP_{sim} was within the uncertainty range of GPP_{EC} are plotted in Figure 3 and are summarized in Table E1.

Table E1

Mean Values of Fractions of Site–Years per Site for Which Annual Average of Simulated Gross Primary Production (GPP_{sim}) Was Within the Uncertainty Range of Eddy-Covariance Derived GPP (GPP_{EC})

		Parameterization strategies						
Models	Per site-year	Per site using Cost ^{IAV}	Per site	Per PFT	Global			
P^W_{hr}	0.869	0.757	0.773	0.481	0.409			
P _{hr}	0.123	0.128	0.131	0.129	0.131			
Bao _{hr}	0.931	0.831	0.840	0.531	0.430			
Bao _{dd}	0.942	0.835	0.838	0.507	0.389			



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Data Availability Statement

The codes that were used to perform all the necessary analyses and plot all the figures in this study are available at https://doi.org/10.5281/zenodo.13729514 (De, 2025). The data from eddy covariance sites are available through FLUXNET at https://fluxnet.org/data/fluxnet2015-dataset/, and the Digital Object Identifier (DOI) of the data set for each site are listed in the Table S8 of Supporting Information S1 (Pastorello et al., 2020; FLUXNET. org, 2024a). The FluxnetEO MODIS version 2 data set is available at https://doi.org/10.18160/0KWD-3RRW (Walther et al., 2022, 2023). The ERA5 data set is available at https://doi.org/10.24381/cds.adbb2d47 Hersbach et al. (2023). The ERA-Interim v2.0 data set has been currently archived by ECMWF; however, it can be accessed following these instructions (Berrisford et al., 2011). Atmospheric CO₂ measurements at Mauna Loa observatory are available at https://doi.org/10.15138/yaf1-bk21 (Thoning et al., 2021).

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